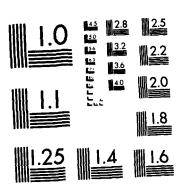
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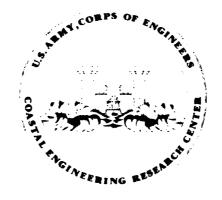


Regional Geology of the Southern Lake Erie (Ohio) Bottom: A Seismic Reflection and Vibracore Study

by

Charles H. Carter, S. Jeffress Williams, Jonathan A. Fuller, and Edward P. Meisburger

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The southern part of the Ohio waters of Lake Erie between Conneaut and			
Marblehead was surveyed in gust of 1977 and 1978 to acquire knowledge of the			
nature, distribution, and geometry of the lake deposits. Primary data consist			
of 576 kilometers of seismic reflection trackline profiles and 58 vibracores			
ranging from 0.7 to 6.1 meters long. About 23 percent of Ohio's part of Lake Erie was covered by the survey.			
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Devonian shale overlain by Quaternary glacial tills and postglacial deposits underlies most of the survey area. In general, the shale is exposed nearest the shore and is succeeded offshore by till and postglacial deposits. The shale surface is commonly more irregular than the till and postglacial deposit surfaces; slopes on the lakeward dipping shale surface range from about 5 to 20 meters per kilometer.

Pleistocene tills—both basal and flow tills—also underlie most of the survey area with extensive exposures between Fairport Harbor and Avon Lake and off Lorain. Interlaminated silts and clays are interbedded with the flow till in some cores; three cores contain both basal and flow tills. The tills are made up largely of silt and clay—size particles composed of quartz and illite. The till has a flatter and more uniform surface than that of the underlying shale, with lakeward slopes ranging from about 1 to 4 meters per kilometer. The till varies in thickness from 0 to 26 meters and thickens lakeward at about 5 meters per kilometer.

Sand, muddy sand, sandy mud, and mud are the four principal postglacial deposits. These deposits commonly lie lakeward and overlie rock and till. In general, the coarser deposits lie nearest the shore. However, the two principal sand deposits at Fairport Harbor and Lorain-Vermilion are well offshore. Also, the finer deposits are found closer to shore and in shallower water west of Cleveland. Combined postglacial deposit thicknesses range from 0 to 22 meters and like the till, the postglacial sediment thickens lakeward.

The tills were first deposited on an irregular, erosional shale surface. Till deposition continued intermittently on both shale and previously deposited till until eastward retreat of the last Wisconsinan glacier from the Erie basin. Drainage of the lake ponded west of the glacier then exposed the till to subaerial erosion which led to the formation of stream channels in the till off Lorain and Fairport Harbor. Isostatic rebound of the outlet then led to a rise in lake level with associated erosion and deposition along the expanding lakeshore, which tended to smooth the till surface. The early postglacial (Holocene) deposits, which accumulated during the rise in lake level and cover the underlying till and shale, were deposited in a complex of fluvial, deltaic, and lacustrine environments. Modern lacustrine muds are now being deposited over these early Holocene deposits.

PREFACE

This report is one of three reports which describe results of the Inner Continental Shelf Sediment and Structure (ICONS) study of southern Lake Erie. The first report (Williams, et al., 1980) deals with the sand resources in Ohio and the second report (Williams and Meisburger, 1982) provides survey results from Pennsylvania. The primary objective of the ICONS program is to locate and delineate sand and gravel deposits suitable for beach nourishment and restoration (Duane, 1968). The work was carried out under the U.S. Army Coastal Engineering Research Center's (CERC) Barrier Island Sedimentation Studies work unit, Shore Protection and Restoration Program, Coastal Engineering Area of Civil Works Research and Development, in cooperation with the Ohio Department of Natural Resources, Division of Geological Survey (DGS).

The report was prepared by Charles H. Carter and Jonathan A. Fuller, Geologists, under the general supervision of H.R. Collins, Chief, DGS, and by S. Jeffress Williams and Edward P. Meisburger, Geologists, under the general supervision of Dr. C.H. Everts, Chief, Engineering Geology Branch, and Mr. N. Parker, Chief, Engineering Development Division, CERC. Data collection was conducted by CERC and DGS with the assistance of U.S. Army Engineer Districts, Buffalo and Mobile, and the U.S. Army Engineer Waterways Experiment Station (WES).

The authors acknowledge the assistance of a large number of people who contributed to the success of this study. J. May, J. Forbes, and D. Andrews of WES operated the seismic reflection equipment; E. Lagrone of the Mobile District operated the vibracore equipment; and M. Chambers of the Buffalo District skippered the tug and scow for the vibracore operation. Within DGS, D.L. Liebenthal skippered the boat carrying the seismic reflection equipment and the navigation system; D.E. Guy, Jr., C.L. Hopfinger, T.J. Feldkamp, J.D. Reed, and J. Vormelker positioned the transponders for the Mini Ranger; R.W. Carlton provided the X-ray diffraction analyses; and C.L. Hopfinger helped throughout with the laboratory work on the vibracores, with data compilation in the office, and with identification of the mollusks, with the aid of M.J. Camp of the University of Toledo. N.A. Rukavina (Environment Canada) provided helpful comments on parts of the report. In addition, G.P. Hall and J.C. Dixon of the Ohio Department of Transportation provided the Atterberg limits, and T.L. Lewis of Cleveland State University had the radiocarbon work done. D.A. Prins of CERC served as field survey chief during both data collection phases, and D.J. Benson, formerly of DGS, helped plan the field surveys and took part in the seismic reflection survey. Lastly, C.H. Everts and H.R. Collins made constructive reviews and their comments are appreciated.

Original copies of all seismic data are stored at CERC. Cores collected during the field survey program in Ohio are in a repository at the University of Toledo, under agreement with CERC. Requests for information relative to these items should be directed to CERC or the Department of Geology, University of Toledo.

Technical Director of CERC was Dr. Robert W. Whalin, P.E., upon publication of this report.

Comments on this publication are invited.

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CONVERSION FACTORS, U.S. CUSTOMARY TO METRIC (SI) UNITS OF MEASUREMENT U.S. customary units of measurement used in this report can be converted to metric (SI) units as follows:

Multiply	by	To obtain
inches	25.4	millimeters
	2.54	centimeters
square inches	6.452	square centimeters
cubic inches	16.39	cubic centimeters
feet	30.48	centimeters
	0.3048	meters
square feet	0.0929	square meters
cubic feet	0.0283	cubic meters
yards	0.9144	meters
square yards	0.836	square meters
cubic yards	0.7646	cubic meters
miles	1.6093	kilometers
square miles	259.0	hectares
knots	1.852	kilometers per hour
acres	0.4047	hectares
foot-pounds	1.3558	newton meters
millibars	1.0197×10^{-3}	kilograms per square centimeter
ounces	28.35	grams
pounds	453.6	grams
pourius	0.4536	kilograms
	0.4330	NATOGE GEO
ton, long	1.0160	metric tons
ton, short	0.9072	metric tons
degrees (angle)	0.01745	radians
Fahrenheit degrees	5/9	Celsius degrees or Kelvins ¹

¹To obtain Celsius (C) temperature readings from Fahrenheit (F) readings, use formula: C = (5/9) (F -32).

To obtain Kelvin (K) readings, use formula: K = (5/9) (F -32) + 273.15.

REGIONAL GEOLOGY OF THE SOUTHERN LAKE ERIE (OHIO) BOTTOM: A SEISMIC REFLECTION AND VIBRACORE STUDY

bу

Charles H. Carter, S. Jeffress Williams, Jonathan A. Fuller, and Edward P. Meisburger

I. INTRODUCTION

1. Background and Scope.

The construction, improvement, and periodic maintenance of beaches and dunes by the placement (nourishment) of suitable sand along the shoreline is an important means of counteracting coastal erosion and enhancing recreational facilities. However, it has become increasingly difficult in recent years to obtain suitable sand from traditional sources such as lagoons and inland deposits because of economic and ecological factors. This problem led the U.S. Army Coastal Engineering Research Center (CERC) to initiate a search for offshore sand deposits. Exploratory efforts to locate and inventory deposits suitable for future fill requirements began in 1964 with a survey off the New Jersey coast (Duane, 1969). Subsequent data collection surveys have included the Inner Continental Shelf areas off New England, Long Island, Delaware, Maryland, Virginia, Florida, the Cape Fear region of North Carolina, eastern Lake Michigan, and southern California. This program, formerly known as the Sand Inventory Program, is now known as the Inner Continental Shelf Sediment and Structure (ICONS) program.

The type of data collected for the ICONS program is not only useful in locating potential borrow areas but is of further value in providing geological information for planning, design, and environmental impact evaluation of coastal engineering works. The results of the ICONS studies are normally presented in two separate but complementary reports: one covering sand resources, the other covering the regional geology.

This study is unique from the previous ICONS studies in that it was conducted in cooperation with a State geological survey, the Ohio Department of Natural Resources, Division of Geological Survey (DGS). The report deals primarily with the bottom and subbottom deposits in the Ohio waters of Lake Erie as mapped largely from high-frequency seismic reflection profiles and vibracores between Conneaut (at the Ohio-Pennsylvania border) and Marblehead. The report does not include the reach between Marblehead and Toledo. Seismic profiling of this reach was done, but no vibracores were taken and the seismic records are difficult to interpret because of the lack of well-defined reflectors. The lack of well-defined reflectors on the records between Vermilion and Marblehead also precluded mapping of the subbottom deposits along this stretch of shore. This report follows a report on sand resources in Ohio (Williams, et al., 1980) and another report on the Pennsylvania part of Lake Erie (Williams and Meisburger, 1982).

The study area encompasses a zone ranging from 1 to 16 kilometers offshore between Conneaut and Marblehead (Fig. 1). Survey coverage of the area is shown

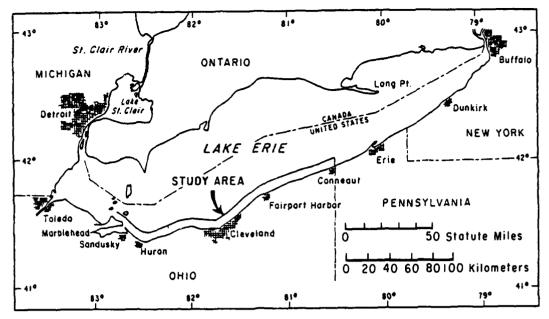


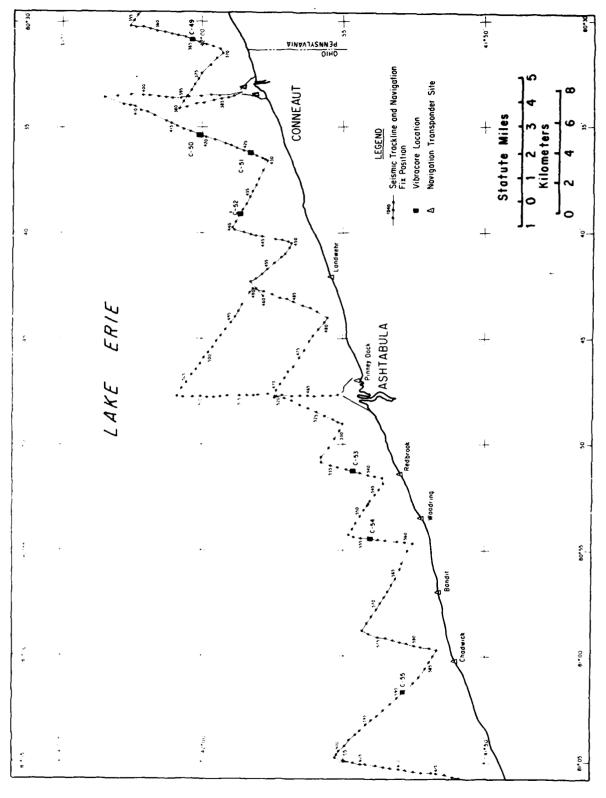
Figure 1. Lake Erie study area, from Conneaut to Marblehead, Ohio.

in Figures 2 to 7. Data collected for this study consist of 576 kilometers of seismic reflection trackline profiles, taken in August 1977, and 58 vibracores ranging from 0.7 to 6.1 meters long, taken in August 1978. About 23 percent of Ohio's open lake part of Lake Erie (7481 square kilometers) was covered by the seismic reflection survey between Conneaut and Marblehead.

The survey data in this report were supplemented in places by previous DGS work. Vertical control was obtained from National Ocean Survey (NOS) water level gage data for Lake Erie; water depths are referenced to low water datum (LWD), which is 173.3 meters above mean water level at Father Point, Quebec (International Great Lakes Datum (IGLD), 1955). Mean lake level in both August 1977 and August 1978 was about 1 meter above LWD. This study is basically reconnaissance in nature, as seismic line spacing and orientation and core spacing density preclude a detailed evaluation of the bottom and subbottom deposits. However, because of the relatively flat-lying nature of the deposits, extrapolation between tracklines can be made with a fair degree of confidence.

2. Field Procedures.

a. Geographic Positioning System. A radar-type electronic positioning system, the Motorola Mini-Ranger III, was used to determine position of the research survey vessel during the seismic survey (phase I) and the vibracoring (phase II). The system determines the position of the survey vessel with respect to two known reference points on shore and is restricted to line-of-sight operation. The basic system consists of a master mobile unit mounted onboard the vessel and two shore-based transponders. The master unit triggers reply pulses from the transponders; each transponder pulse is received separately and the elapsed time between the transmitted pulse and the individual transponder reply pulse is converted to a measurement of distance. Each distance (range) from the two transponders at the known shore stations is displayed in turn on the range console. This range information, together with the known locations of the shore stations, is then trilaterated and plotted on hydrographic charts to obtain the position (fix) of the survey vessel.



Trackline location map, Ohio-Pennsylvania State line to east of Fairport Harbor. Figure 2.

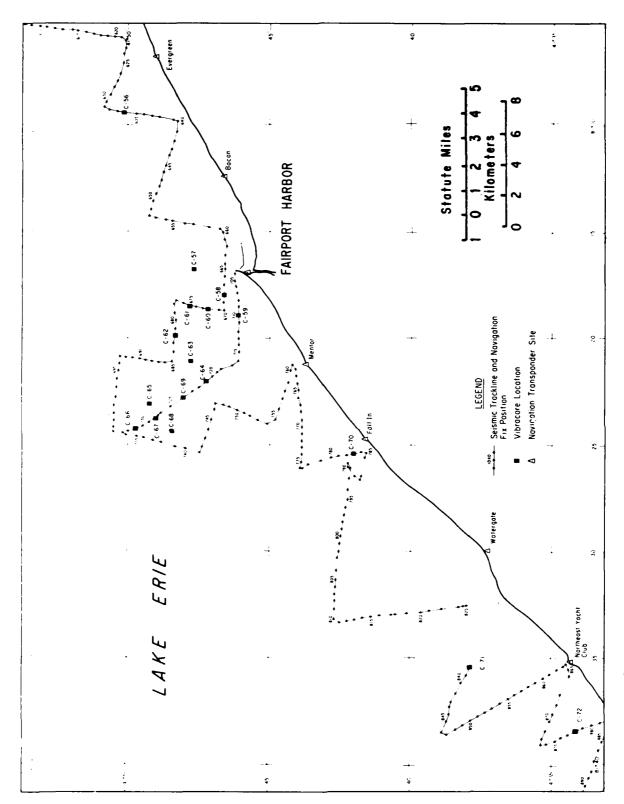
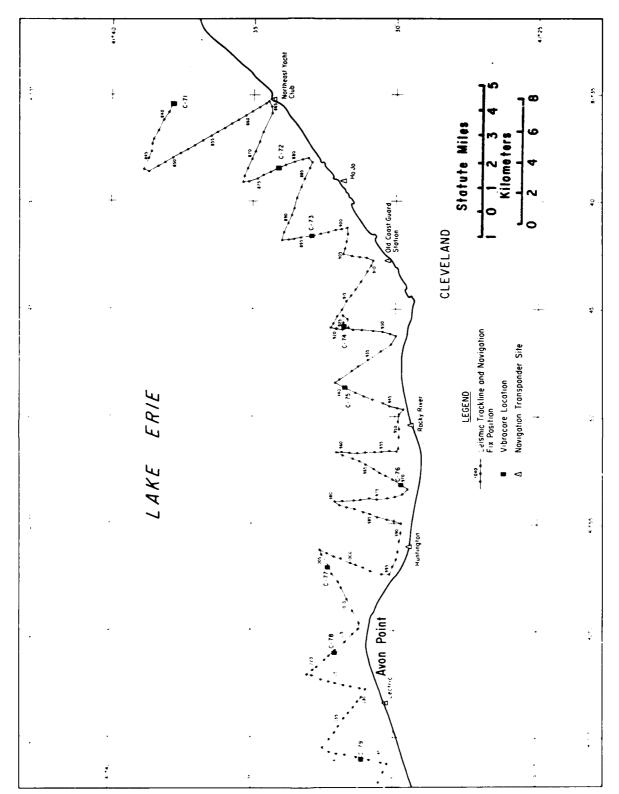


Figure 3. Trackline location map, east of Fairport Harbor to Northeast Yacht Club.



Trackline location map, Northeast Yacht Club to Avon Point. Figure 4.

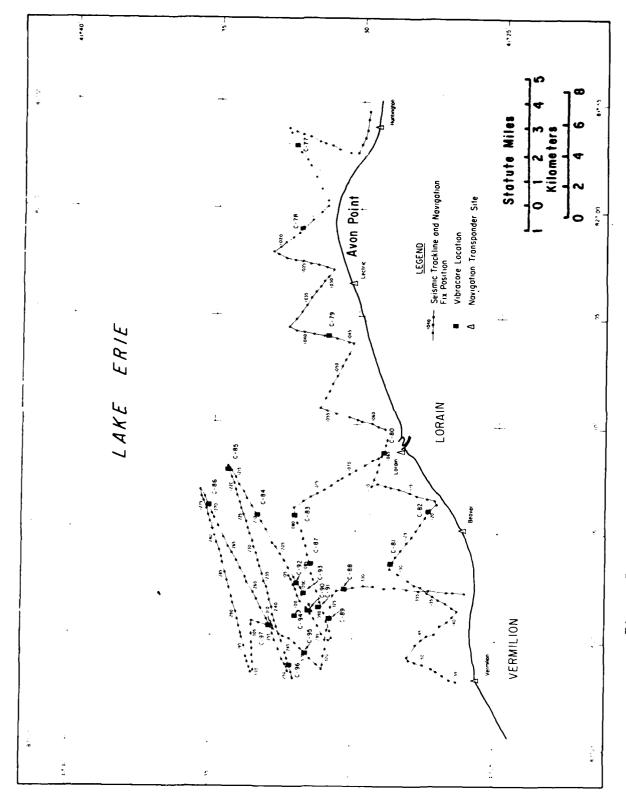


Figure 5. Trackline location map, Avon Point to Vermilion.

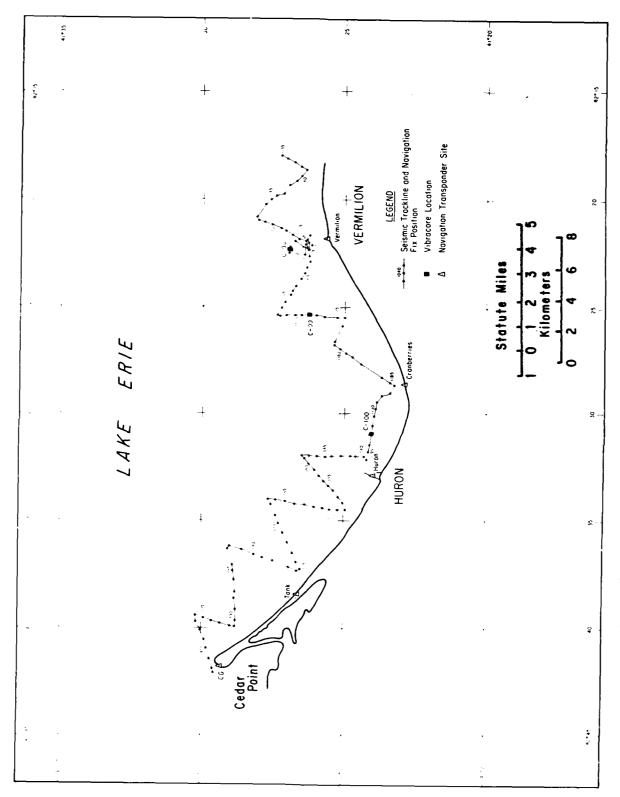


Figure 6. Trackline location map, Vermilion to Cedar Point.

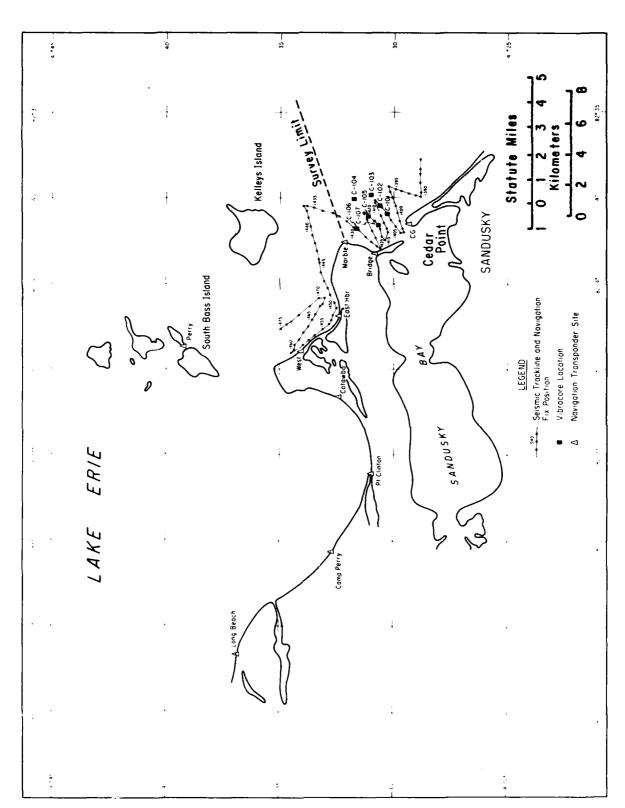


Figure 7. Trackline location map, Cedar Point to Long Beach.

Navigational fixes during the seismic survey were obtained about every 2 minutes and each fix was keyed to the seismic records by an event mark on the records.

b. Seismic Reflection Profiling. Seismic reflection profiling is widely used to delineate geologic features on land and over water. In this study, repetitive high-energy sound pulses were generated near the water surface to produce seismic waves, which reflected off the geologic features; these waves were received and recorded on paper by a recorder aboard the survey vessel to produce a cross section representing the features at and below the lake bottom.

The seismic reflection data were obtained by towing the sound-generating and receiving instruments behind the DGS research vessel, GS-1 (Fig. 8), which followed predetermined survey tracklines (Figs. 2 to 7). In phase I of this study, two seismic subbottom profiling systems were used simultaneously: an Ocean Research Equipment, Inc. (ORE) 3.5-kilohertz pinger system and an Edgerton, Gremerhausen and Greer (EG&G), Inc. UNIBOOM system. The pinger records were of little use because of their lack of resolvable reflectors, whereas the UNIBOOM records were generally good. The inshore boundary of the survey was about 1 kilometer from shore and the offshore boundary was about 7 kilometers offshore. The nearshore survey water depths were about 7.5 meters, which is about the minimum depth for obtaining good-quality seismic profiles; the offshore water depths were about 14 meters. Information on various seismic profiling techniques is discussed in Ewing (1963), Moore and Palmer (1967), Barnes, et al. (1972), and American Association of Petroleum Geologists (1977).



Figure 8. The DGS Research Vessel GS-1 used to tow the seismic equipment and locate core sites.

Coring Equipment. A pneumatic vibratory coring device designed to obtain continuous sediment cores a maximum of 6.1 meters long was used in the phase II survey operation (Fig. 9). The apparatus is equally effective in penetrating and recovering granular and cohesive sediments; however, a stony till is not easily penetrated and the core barrel will not penetrate coherent rock. The core rig consists of a 10-centimeter steel core barrel, clear plastic inner liner, shoe and core catcher, and a pneumatic driving head attached to the upper end of the barrel. These elements are enclosed in a quadrapodlike frame with four articulated legs which rest on the lake bottom. An aluminum H-beam and frame serve as a support structure and guide for the vibrator head and core pipe as the core barrel penetrates the lake bottom. The lack of rigid attachment of the coring device to the surface vessel allowed limited motion of the vessel during the actual coring processes. Power was supplied to the pneumatic vibrator head by a flexible hoseline connected to a large-capacity (118 liters per second) air compressor. After coring was completed, the assembly was hoisted onboard the vessel, the liner containing the core was removed, samples from the top and bottom of the core recovered, the ends sealed, and the core carefully marked for orientation and identification. The historical development of vibratory coring equipment is discussed by Tirey (1972).

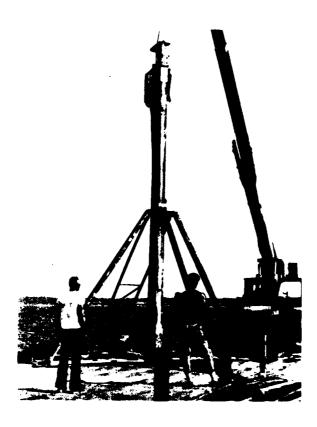


Figure 9. A 6-meter-long vibratory coring apparatus used to collect sediment cores is shown being lifted off the platform for deployment on the lake floor.

A 36-meter-long scow from the U.S. Army Engineer District, Buffalo, was used as the platform for phase II coring. The scow was transported by the Corps tugboat *Washington* (Fig. 10).

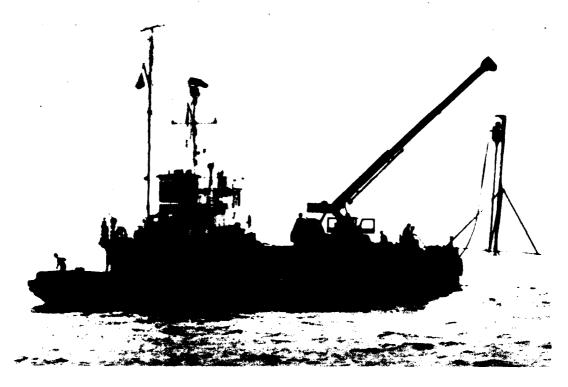


Figure 10. The U.S. Army Engineer District, Buffalo, tugboat Washington.

d. <u>Data Collection</u>. Before the fieldwork began, seismic survey tracklines were plotted on navigation charts of the survey area. Position, spacing, and length of the tracklines were determined by several factors. The primary concern was spacing the lines to achieve maximum coverage of the study area within the limits of time and budget. After the survey tracklines were selected, the locations of the shore stations for the navigation system were determined. Of high priority were stations at elevated positions (for adequate line-of-sight), which also offered good triangular position in relation to the survey vessel and adjacent shore stations. (Acceptable results are achieved when the angle of range intercept of the vessel is greater than 30° and less than 150°; optimum range angle intercept is 90°.) A total of 32 shore navigation stations were used along 207 kilometers of coast from Conneaut to Marblehead. Positions and spacing of the tracklines were altered at times to gather additional information on geologic features such as buried stream channels, sediment contacts, and probable lake bottom exposures of sand.

Interpretations of the seismic profile records were made to select coring sites with the greatest potential for sand and subsurface information. These records were visually examined and marked to establish the primary geologic features (e.g., regional sedimentary reflectors, sediment contacts, buried stream channels). Selected acoustic reflectors were then mapped to provide areal continuity of horizons considered significant because of their areal extent and relationship to the general structure and geology of the study area. The use of seismic data to interpret geologic conditions before selecting the core sites maximizes the usefulness of both sources of data and provides the most efficient use of funds.

During phase II, the vessel GS-1 was used to relocate fix positions selected as coring sites by duplicating the range values from the shore stations. The GS-1 first maneuvered until one of the ranges was duplicated and then an arc was run on that range until the other range was intersected, at which time an anchored float was used to mark the core location. Core sites were located and marked in this manner because of the limited maneuverability of the scow. The GS-1 crew located a core position in minutes and dropped a float marker; the tug and scow then moved in on the marker, anchored, and the core rig was lifted from the deck of the scow and set on the lake bottom. Meanwhile, the GS-1 proceeded to the next core site. Once the core rig was on the bottom, the core barrel was driven into the lake bottom sediments; within about 15 minutes the coring was completed and the apparatus was lifted back onto the scow. The core liner containing the sediment was removed from the barrel and small reference samples were obtained from the top and the bottom of each core. The liner was then capped and sealed, labeled, and a general description of the samples was made. The scow was then moved to the next coring location. While underway, the corer was reassembled and loaded with a new liner. In general, the corer penetrated the lake bottom deposits quite easily; however, penetration in stony till and in some well-sorted, medium-grained sands was poor.

3. Office and Laboratory Procedures.

After completion of the data collection, the navigational fix marks, ship trackline positions, core sites, and shore stations were plotted to show the coverage in the survey area (Figs. 2 to 7). The cores were split longitudinally, using a circular powersaw to cut the plastic liner and a piece of thin wire to cut the sediment. As soon as the core was opened it was color typed using the Munsell color system (Munsell Soil Color Charts, 1954 ed., Munsell Color, Inc., Baltimore, Md.). Color typing was immediately followed by sampling for natural water determinations (sample is weighed, dried, and reweighed), by unconfined compressive and shear-strength measurements (using hand-held Soiltest instruments), and by detailed descriptions (see App. A). The unsampled half of the core was wrapped in plastic for storage. After the sampled half had partially dried (a day or two), the description was checked, photos of representative units were taken, and additional sampling was done if necessary.

Three principal techniques were used for the grain-size analyses: rapid sand analysis (RSA) for the finer sands (U.S. Army, Corps of Engineers, Coastal Engineering Research Center, 1977), sieve analysis at 0.5-phi intervals for the coarser sands and fine gravel (Folk, 1974), and pipet analysis for the samples containing appreciable clay- and silt-size particles (Folk, 1974) (App. B). In addition, 22 clay samples were analyzed by X-ray diffraction (App. C), and general sand-grain mineralogy was determined with a binocular microscope using a feldspar stain technique (Gross and Moran, 1970). Atterberg limits were determined for 29 fine-grained samples (App. D), and two radiocarbon-14 ages were determined from wood samples in core 62. Mollusks from selected samples were identified generally to species level (App. E).

4. Seismic Interpretation.

Seismic reflections from the shale and till surfaces were extensive enough in the subbottom to be mapped throughout most of the study area. The shale reflector generally is broken and irregular but well defined with closely

spaced internal reflectors which either intersect the main reflector at a low angle or parallel it in a regular pattern (Fig. 11). On the other hand, the till reflector generally is continuous and not as irregular nor as well defined as the shale reflector. In addition, irregular but fairly continuous internal reflectors are fairly common in the till (Fig. 12). The shale and till reflectors were mapped at two confidence levels. The first confidence level was used where the reflector was well defined on the seismic record. The second confidence level was used where the reflector was not as well defined but its position could be reasonably well inferred from the overall geologic setting and character of the reflector. In addition, the echo character type of the surface sediment, in conjunction with reference sediment samples, was used to map the surface sediment. Seven surface echo character types were defined: rock, till, sand, muddy sand, sandy mud, mud, and rock waste (Table 1).

Following the mapping of the reflectors, acoustic traveltimes were neasured to determine sediment thicknesses (App. F). The velocities of sound through the till and postglacial sediment are used to convert these measurements into sediment thickness. The velocity of sound in water was calculated to be about 1.54 kilometers per second from measured depths at several core locations (App. G). The average velocity in the postglacial sediment was calculated to be about 1.3 kilometers per second in a similar way by using the thickness of the postglacial sediment in the vibracores (Apps. H and I). A velocity for the till was not calculated because none of the cores penetrated till to rock; however, laboratory work by Morgan (1964) suggests that 1.8 kilometers per second is a realistic velocity for Lake Erie till so this value was used in the calculations.

5. Previous Studies.

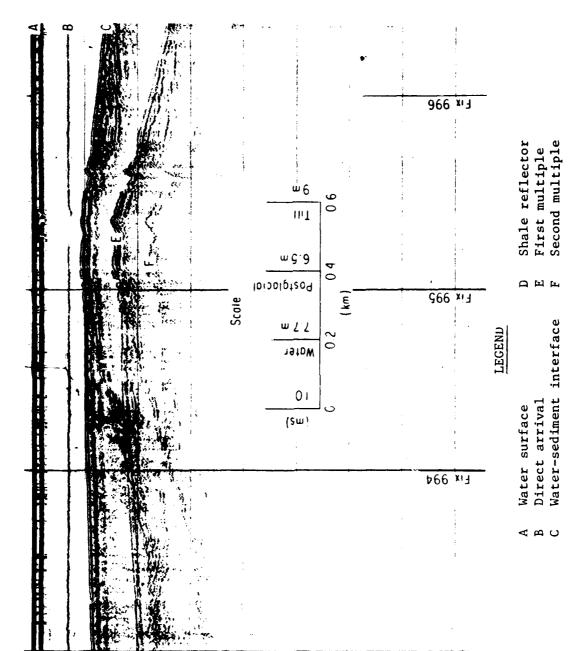
The first comprehensive bottom sediment study of the Ohio part of the central Lake Erie basin was done by Hartley (1961). Bottom grab samples were taken on 1.6- or 3.2-kilometer grids and some subbottom sampling was done by coring or jetting. Hartley's work provided important data that were used extensively in the planning stages of this study. Shore and nearshore deposits within 610 meters of the shoreline were mapped in the 1970's as part of DGS's county shore erosion studies (Benson, 1978; Carter and Guy, 1980); this information was used in the current study to map the sediment adjacent to the shore. Also, several hundred kilometers of seismic reflection profiles and 32 borings and vibracores were collected in a rectangular nearshore area off Cleveland (Dames and Moore, 1974).

Three shallow seismic reflection surveys of central Lake Erie have been reported (Morgan, 1964; Lewis, 1966; Wall, 1968). The works by Morgan (1964) and Wall (1968) are too general to be of use, whereas the work by Lewis (1966) is sufficiently detailed, particularly on the geologic history of the lake, and of use.

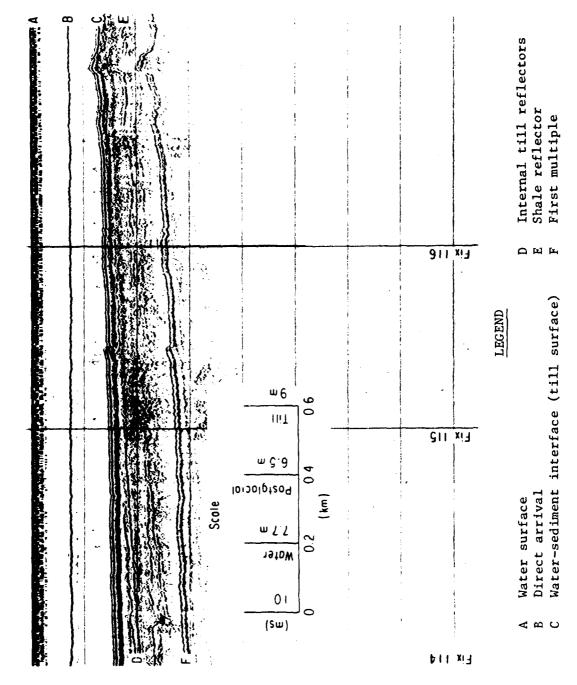
II. PHYSICAL SETTING

1. Introduction.

The Lake Erie basin is underlain by middle Paleozoic sedimentary rocks that are overlain by Pleistocene- and Holocene-age deposits. The Pleistocene deposits



Seismic record showing character of shale reflector; this section is about 3 kilometers east of Avon Point. Figure 11.



Seismic record showing internal reflectors in till deposit; this section is about 2 kilometers west of Lorain. Figure 12.

Table 1. Surface-sediment echo character and their interpretation.

Echo character	Sediment interpretation	Letter designation
Very poor multiple, l smooth surface, mostly transparent, may have continuous, multiple, parallel internal reflectors.	Soft, nearly fluid mud	М
Very poor to poor multiples, last slightly irregular surface (rougher than M), nontransparent (a noisy record).	Sandy mud	Х
Good multiples (commonly two to three), smooth to slightly irregular surface (similar to X), nontransparent.	Muddy sand	0
Fair multiples, hummocky to gently rolling but smooth surface, commonly nontransparent.	Sand	S
Poor to good multiples (mostly fair), smooth or irregular surface (in between S and R), commonly irregular internal reflectors.	Clay and clay with gravel (probably till)	Т
Fair to good multiples (commonly three or more), commonly a broken, irregular surface (may look similar to T), and may have closely spaced discontinuous, internal reflectors.	Rock (probably shale)	R
Fair to good multiples and a jagged, irregular surface.	Rock waste (shale derived from a harbor excavation)	R W

¹Very poor multiple: a discontinuous multiple if present.

Poor multiples: one to two multiples; the first multiple is commonly discontinuous.

Fair multiples: one to two multiples; the first multiple is commonly continuous.

 ${\tt Good\ multiples:}\ two\ to\ seven\ multiples;$ the first and second multiples are commonly continuous.

consist largely of glacial till and the Holocene deposits consist largely of fine-grained lake sediments. About 14,000 years before present (B.P.) a glacial lake occupied the Lake Erie basin to an elevation of about 244 meters above sea level (Sly, 1976). Within the next 1,400 years or so, lake level fluctuations of a few tens of meters took place as the receding glacier, in a series of retreats and advances, alternately exposed and buried different lake outlets. The glacial activity ended about 12,600 years B.P. with retreat of the last Wisconsinan glacier. This allowed the lake waters to discharge northeastward across the Niagara escarpment, which at that time was depressed due to the weight of the glacier. Because this outlet was about 40 meters below present lake level, the level of early Lake Erie was at about 134 meters above sea level. Subsequent isostatic rebound of the escarpment at the outlet led to the filling of the lake to its present elevation of about 174 meters above sea level.

Shore.

The shore from Conneaut to Huron consists of moderate to high relief (3 to 20 meters) shale and till slopes and bluffs commonly capped by old lake deposits. The shale is exposed for appreciable lengths above or near lake level from Conneaut to Fairport Harbor, near Euclid (Moss Point), between Cleveland and Avon Point, and near Vermilion. The shore from Huron to Marblehead consists of low relief (<2 meters) barrier beaches and laminated clay banks. The shore at Marblehead consists largely of moderate relief (3 to 6 meters) dolostone and limestone banks.

Beaches make up a fragmented band that front about 50 percent of the shore in the study area. The beaches are commonly narrow (<15 meters wide) and consist primarily of sand, although there are pocket beaches of cobbles in places where the shore is composed of rock. In addition, there are about 3,000 shore protection structures consisting largely of groins and seawalls scattered along the shore.

Offshore.

Lake Erie is the shallowest of the five Great Lakes with an average depth of about 19 meters and a maximum depth of 65 meters. The nearshore slopes are generally less than 1°; in the study area within 1 to 2 kilometers of the shore the slope is about 0.25°, except between Fairport Harbor and Avon Point, where the slope is about 0.50°. Offshore the slopes become gentler with a slight decrease in slope from east to west; the area off Lorain-Vermilion is the principal exception. In this area the offshore bottom rises to form a subtle ridge (Pelee-Lorain ridge) which extends across the lake to Point Pelee, Canada. Sand and gravel, generally less than 2 meters thick, cover much of the nearshore (<1 kilometer offshore) bottom; most of the bottom farther offshore is made up of fine-grained postglacial sediment, rock, till, or glaciolacustrine clay (Verber, 1957; Hartley, 1961; Williams, et al., 1980). In general, the distribution of these offshore deposits (to the international boundary) can be divided into two areas: from Conneaut to Fairport Harbor, a band of shale is bordered offshore by sand and finer postglacial sediment; from Fairport Harbor to Marblehead, a band of sand and gravel, till, or rock (mostly shale) is bordered offshore by finer, postglacial sediment except for till in the Cleveland area and sand and gravel in the Lorain-Vermilion area.

III. BOTTOM DEPOSITS

1. Introduction.

Shale, till, and postglacial sediment ranging from sand to mud make up most of the bottom deposits. Shale underlies most of the Ohio part of Lake Erie. The shale is overlain by glacial tills that are overlain by postglacial deposits. In general, the shale is exposed nearest the shore and succeeded offshore by the tills and postglacial deposits.

Shale.

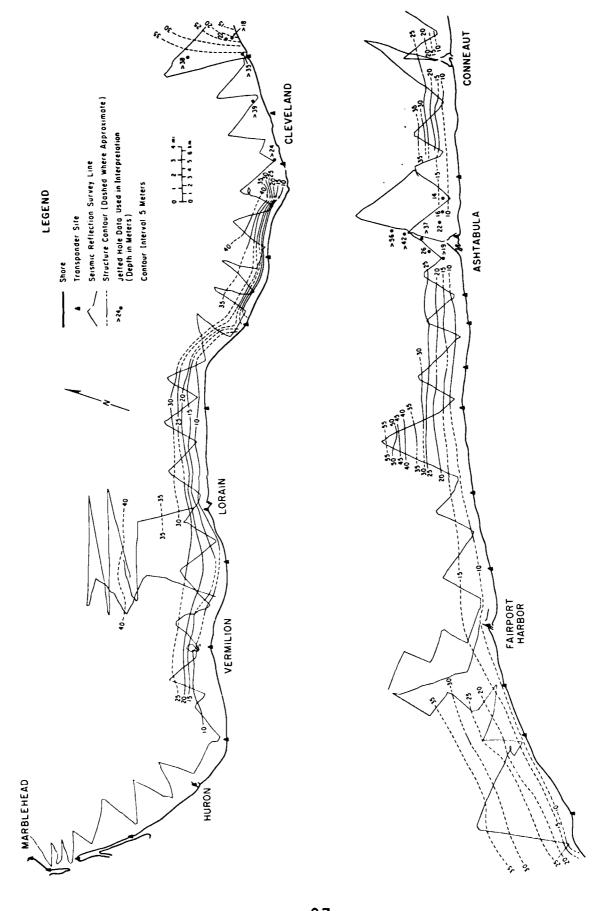
The contact between the Devonian limestones and dolomites to the west and the Devonian shales to the east lies between Huron and Sandusky (Bownocker, 1920). The northeast trend of this contact, a gentle southeast dip, and shore and nearshore exposures indicate that the Devonian Ohio Shale underlies most of the survey area. In northern Ohio, the shale is made up of two carbonaceous, blue-black shale members separated by a largely noncarbonaceous, blue-gray, silty shale member (Hoover, 1960).

Shale is exposed close to the shore in a continuous 2-kilometer-wide band between Conneaut and Fairport Harbor, in a broken 1- to 2-kilometer-wide band between Cleveland and Lorain, and in a continuous 2- to 3-kilometer-wide band off Vermilion. The echo character of the shale surface on the seismic records ranges from smooth to irregular (Fig. 11). In the subsurface, the shale reflector could be mapped along most of the tracklines; however, the reflector was not apparent in places off Ashtabula, Fairport Harbor, Cleveland, and west of Vermilion (Fig. 13).

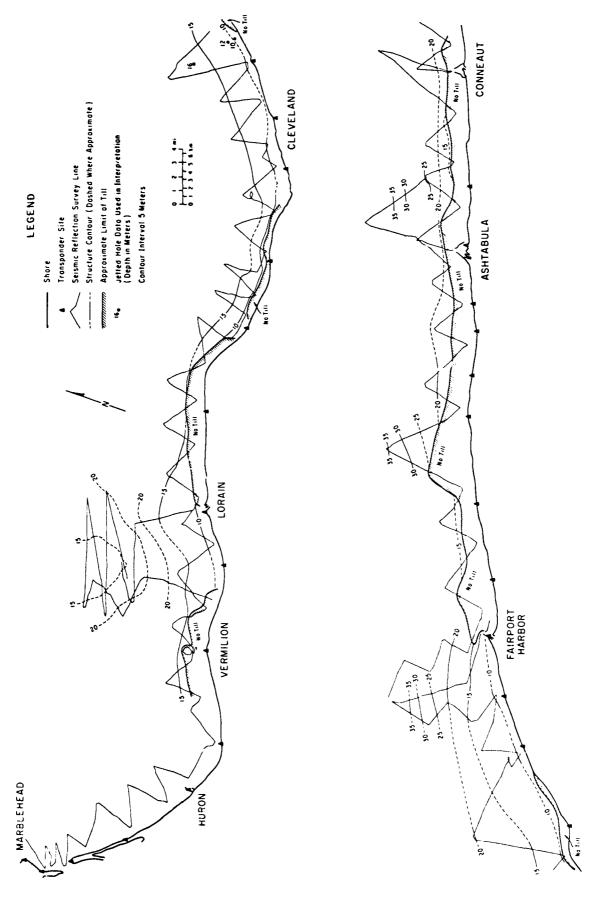
The shale surface slopes lakeward between Conneaut and Ashtabula at about 10 meters per kilometer. West of Ashtabula, the slope is about 5 meters per kilometer, decreasing to about 2 to 3 meters per kilometer at Fairport Harbor and then increasing to about 10 meters per kilometer at Moss Point. The seismic records do not show a shale reflector between Moss Point and west Cleveland. From west Cleveland to Avon Point the slope is 15 to 20 meters per kilometer; from Avon Point to west Vermilion the slope is gentler, about 4 to 6 meters per kilometer. The irregular nearshore shale surface characterized by these variable slopes is further exemplified by a major buried river valley (about 90 meters deep) entering Lake Erie at Cleveland (Crowell, 1979). In addition, there are major valleys eroded into the shale at Huron (Stein, 1962) and Lorain (Pree, 1962).

3. Till.

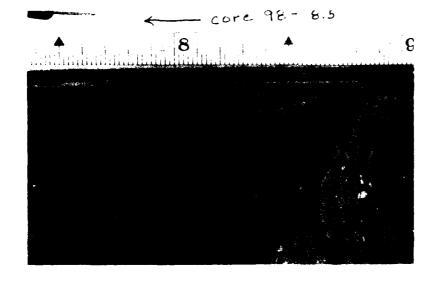
Till overlies the Devonian Ohio Shale over most of the region covered by the tracklines (Fig. 14). It can generally be characterized as a hard, stony, unstratified deposit interpreted as basal till (Fig. 15; gravelly clay in App. A) or a soft, clay-rich, commonly stratified deposit interpreted as flow till (Fig. 16). Some of the till sections (e.g., cores 95, 98, and 100) contain interlaminated silts and clays resembling varves (Fig. 17; clay in App. A); cores 92, 98, and 99 contain both basal and flow till. The tills are made up largely of silt- and clay-size particles composed of quartz and illite. Detailed information on the texture, composition, and engineering properties of the till is in Appendixes B and C; Table 2 is a summary of the information.



Contours Rock structure contour map as interpreted mainly from the seismic reflection records. Contour in meters below survey lake level of about 174.3 meters (IGLD, 1955). The overlying sediment thickness was determined from till and postglacial sediment thicknesses. Figure 13.



Till structure contour map as interpreted mainly from the seismic reflection records, Contours in meters below survey lake level of about 174.3 meters (IGLD, 1955). Figure 14.



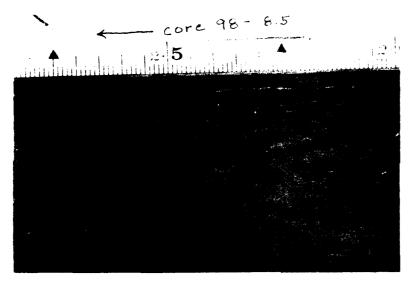


Figure 15. Photo of basal till in core 79; scale in meters from top of core. Arrow points to top of core.

Figure 17. Photo of varved (?) clay in core 98; scale in meters from top of core. Arrow points to top of core.

in meters from top

of core. Arrow points to top of

core.

Photo of flow till in core 98; scale

Figure 16.

Table 2. Properties of basal till and flow till in study area.

	Basal till		Flow till	
Color	Olive gray	Yellowish gray	Olive gray	Variegated
Grain size (weight pct) sand silt clay	22 40 38	16 39 45	7 46 47	13 46 41
Natural (weight pct) water content	16	16	21	20
Unconfined compressive strength (kg/cm ²)	1.8	1.8	0.7	0.9
Shear strength (kg/cm ²)	0.7	0.6	0.3	0.4
Atterberg limits liquid limit plasticity index	26.7 9.1	1 	32.0 11.8	29.2 10.4

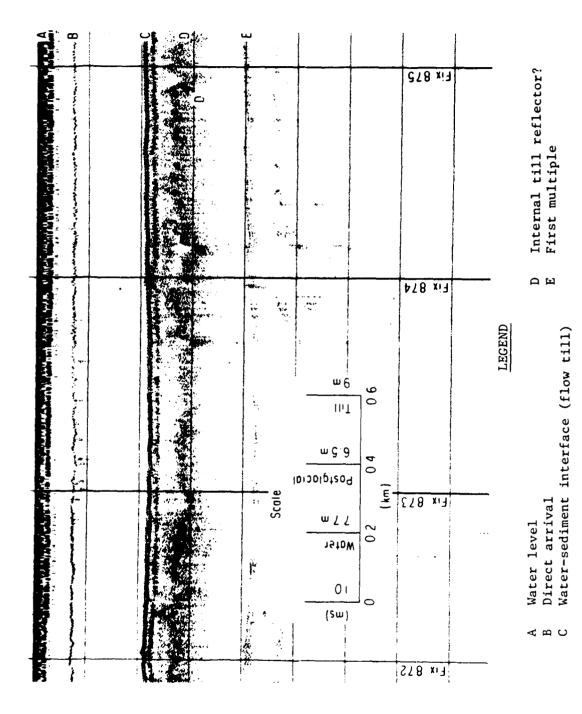
l No data.

Two types of basal till were found in the cores. An olive-gray basal till was found between Ashtabula and Fairport Harbor (cores 54, 55, 60, 61, and 70) and between Avon Point and Huron (cores 77 to 80, 82, 86, 98, and 99); a yellowish-gray basal till was found in the Lorain-Vermilion area (cores 84, 85, 92, 97, and 98). The yellowish-gray basal till overlies the olive-gray basal till in core 98.

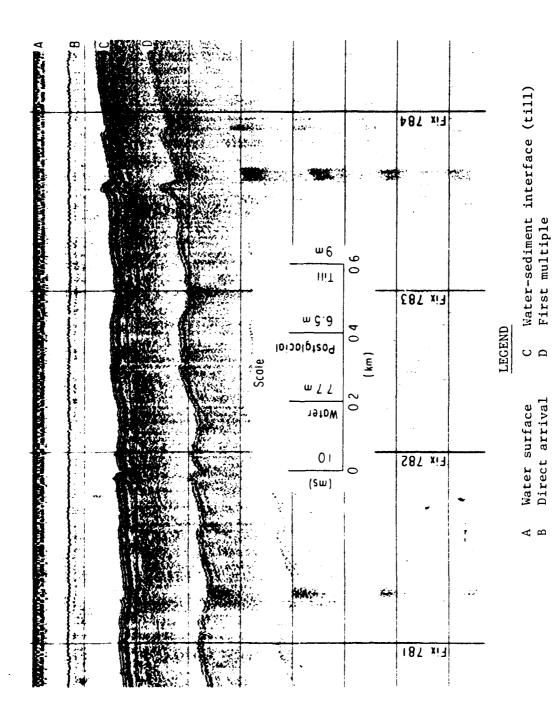
Flow till is found in the Cleveland area and between Lorain and Huron (cores 71 to 76, 81, 82, 92, 95, 96, 98, 99, and 100). A nearly structureless olivegray flow till with minor pea gravel and reddish-brown clay pods was found in the Cleveland area (cores 71 to 76). The flow till between Lorain and Vermilion (cores 81, 92, 95, 96, 98, 99, and 100) has contorted flow banding and ranges in color from olive gray to dark yellowish brown. This flow till overlies a basal till in cores 92, 98, and 99.

The basal till and flow till are exposed nearly continuously between Fairport Harbor and Avon Point and between Lorain and Vermilion. Their surface echo character on the seismic records ranges from smooth to irregular. In general, the till in the Cleveland area has a smooth surface echo character (Fig. 18); the till exposed both east and west has a more irregular surface echo character (Fig. 19).

The till surface from Conneaut to Fairport Harbor slopes lakeward at 2 to 4 meters per kilometer. Between Fairport Harbor and west Cleveland the slope is gentler, 1 to 2 meters per kilometer. From west Cleveland to Avon Point the slope is steeper but fairly uniform at 2 to 3 meters per kilometer.



Seismic record showing the smooth surface echo character of the flow till in the Cleveland area. Figure 18.



Seismic record showing the irregular surface echo character of the till; this section is from about 13 kilometers west of Fairport Harbor. Figure 19.

Direct arrival

Off Lorain-Vermilion the surface slopes lakeward toward the Pelee-Lorain ridge at about 2 meters per kilometer to a depth of 23 meters before rising at the same slope. In general, the till surface is fairly flat in contrast to the underlying shale; however, topographic lows on the till surface were mapped off Lorain-Vermilion (Hartley, 1960) and off Fairport Harbor (Wall, 1968).

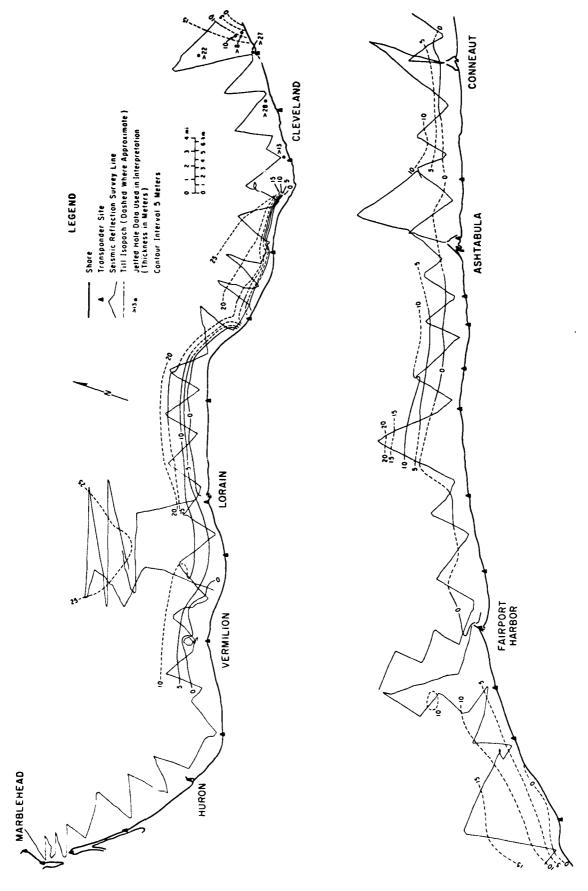
The till thickens offshore, forming a wedge on the underlying rock; till thicknesses determined from the seismic records range from 0 to 26 meters. However, data indicate a till thickness of more than 28 meters off Cleveland (Dames and Moore, 1974). Between Conneaut and Moss Point the till thickens offshore at about 4 meters per kilometer. The till at west Cleveland thickens offshore at about 31 meters per kilometer and then decreases uniformly westward to about 6 meters per kilometer at Lorain (Fig. 20).

Postglacial Sediment.

Postglacial sediments ranging from sand to mud commonly lie lakeward of and overlie rock and till (Fig. 21). The principal exceptions are off Cleveland, where postglacial sediment is succeeded offshore by till, and west of Huron, where postglacial sediment lies adjacent to the shore. The sediments can be classified into four echo character types: sand, muddy sand, sandy mud, and mud (Table 1; Figs. 22 to 25). In addition, rock waste from harbor dredging was mapped off Ashtabula. In general, the sand is less than 0.5 millimeter in diameter and is composed largely of subangular to subrounded grains of quartz, feldspar, and rock fragments. The clay-size particles are similar in mineral-ogy to the clays in the till--mostly quartz, illite, and chlorite--although several samples contain montmorillonite. Macrofossils (clams and snails) were common in some cores, primarily in the coarser sediment. Specific data on the texture, composition, engineering characteristics, and fossils of these deposits are given in Appendixes B to E.

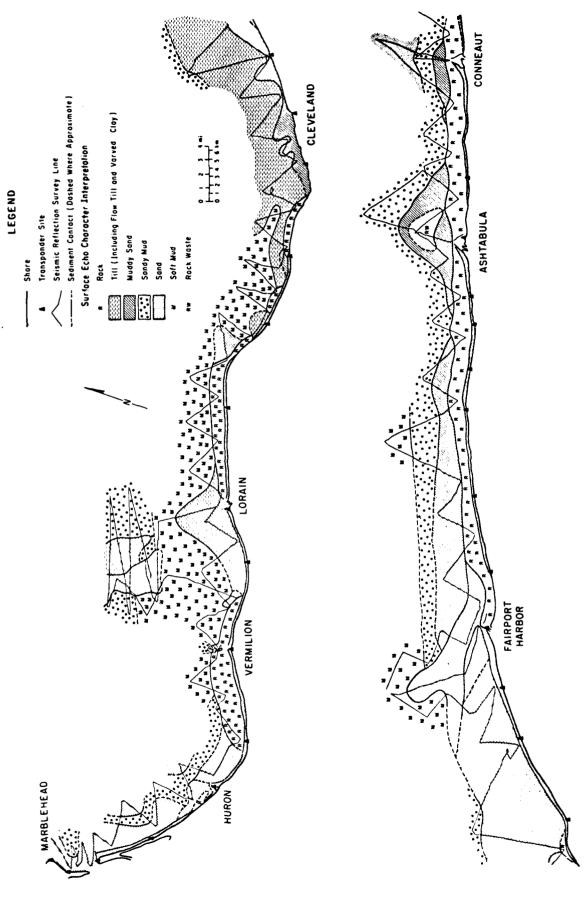
Between Conneaut and Fairport Harbor, muddy sand directly overlying shale is generally found close to shore. This sediment is succeeded offshore by sandy mud, which is commonly succeeded by mud. Between Fairport Harbor and Cleveland, with the exception of the complex Fairport Harbor area, sandy mud was the only postglacial sediment mapped at the lakeward ends of the traverses. Between Cleveland and Vermilion mud generally lies lakeward of till or rock, although coarser grained deposits lie offshore of the mud in the Lorain-Vermilion area. West of Vermilion muddy sand is commonly succeeded lakeward by sandy mud. Sand was mapped in three areas: Conneaut, Fairport Harbor, and Lorain-Vermilion; see Williams, et al. (1980) for a discussion of the sand deposits at Fairport Harbor and Lorain-Vermilion. Also, the finer sediments (mud and sandy mud) are generally found closer to shore and in shallower water west of Cleveland.

The postglacial sediment increases in thickness offshore, except off Lorain, reaching 18 meters between Conneaut and Fairport Harbor, nearly zero between Fairport Harbor and Lorain, and up to 8 meters thick between Lorain and Vermilion (Fig. 26). The thicknesses, however, are difficult to compare because of the short (<5 kilometers) seismic traverses west of Cleveland. Along the Conneaut to Cleveland reach, there is a principal difference at about 9 kilometers offshore; from Conneaut to Fairport Harbor postglacial sediment thicknesses range from 10 to 16 meters, whereas from Fairport Harbor to Cleveland there is an apparent lack of postglacial sediment. Along the



C

Contours in meters Till isopach map as interpreted mainly from seismic reflection records. below survey lake level of about 174.3 meters (IGLD, 1955) Figure 20.



Map of surface sediments as interpreted from the seismic reflection records and core samples. Sediment mapped landward of the tracklines is based on DGS data. Figure 21.

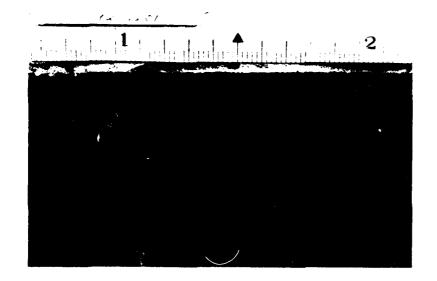


Figure 23. Photo of muddy sand in core 92; scale in meters from top of core. Arrow points to top of core.

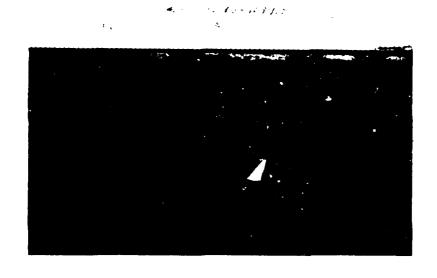


Figure 22. Photo of sand in core 93; scale in meters from top of core. Arrow points to top of core.

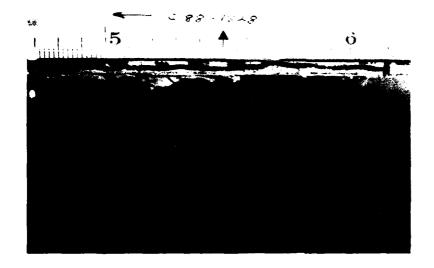


Figure 25. Photo of mud in core 88; scale in meters from top of core. Arrow points to top of core.

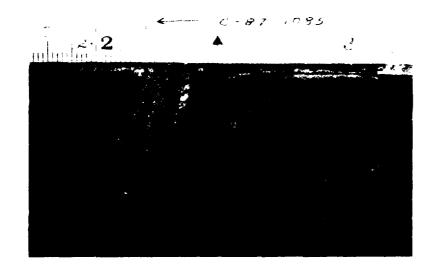
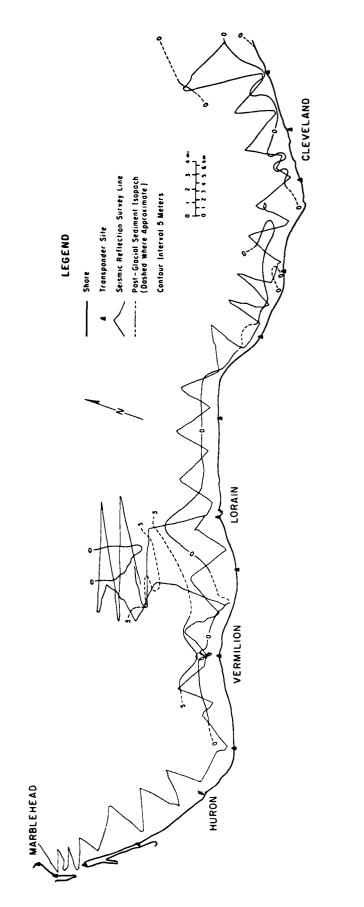
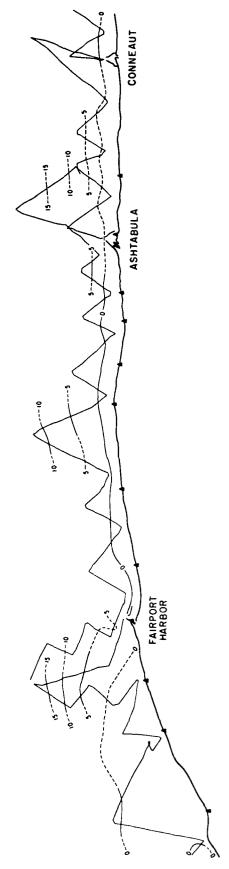


Figure 24. Photo of sandy mud in core 87; scale in meters from top of core. Arrow points to top of core.



C

E



Postglacial sediment isopach as interpreted mainly from seismic reflection records. Contours in meters below survey lake level of about 174.3 meters (IGLD, 1955). Figure 26.

Cleveland to Vermilion reach, with the exception of the area offshore of Lorain, there is also a basic difference at about 5 kilometers offshore; from Cleveland to Lorain postglacial sediment thicknesses range from 0 to 2 meters, whereas from Lorain to Vermilion postglacial sediment thicknesses range from 4 to 5 meters. Offshore of Lorain, the postglacial sediment thickens to 8 meters about 8 kilometers offshore and then decreases to the north as it pinches out against the Pelee-Lorain ridge. Thus it appears that the nature and thickness of the postglacial sediment is largely controlled by the configuration of the underlying rock and till surfaces as well as by the existing wave climate.

IV. GEOLOGIC HISTORY

1. Introduction.

Sedimentary deposits from the Devonian and the Quaternary periods are exposed in the study area. The Devonian Ohio Shale underlies most of the area except for lower Devonian carbonates near Sandusky; Quaternary glacial and interglacial Pleistocene deposits and postglacial Holocene deposits overlie the Devonian rocks. A good general overview of the Quaternary geologic history of the Lake Erie basin is given in Sly and Lewis (1972).

2. Shale.

Seismic reflection profiles extending the width of the central Lake Erie basin illustrate the relatively irregular rock surface on a regional scale (see Fig. 23 in Lewis, 1966). The irregular rock surface, as well as the buried ancestral valleys of the Cuyahoga, Black, and Huron Rivers, indicates that appreciable erosion of the shale has occurred. The relative narrowness and northerly trend of these valleys, in addition to their Pleistocene fill, imply that much of the erosion took place before Pleistocene Glaciation, probably by streams.

3. Till.

The relatively smooth till surface on the seismic records is judged to be the product of several Wisconsinan glacial advances and retreats with glacial and periglacial processes tending to smooth and fill irregularities in the earlier glacial deposits. This hypothesis is consistent with the physical variations in the deposits mapped from the seismic reflection records and the internal reflectors, which indicate that the till is likely made up of multiple till deposits. Field studies on both sides of the lake have documented multiple Wisconsinan tills (Dreimanis, 1970; White and Totten, 1979). Following the classifications of White and Totten (1979) and White (1979), the olive-gray basal till in the cores is interpreted to be the Titusville-Millbrook Till and the yellowish-gray basal till to be either the Lavery-Hayesville Till or the Hiram Till.

The topographic lows in the till mapped off Fairport Harbor (Wall, 1968) and Lorain-Vermilion (Hartley, 1960) are the most prominent features in the till surface. In agreement with Wall's interpretation (1968), the lows represent channels cut into the till surface by eastward-flowing streams probably following the retreat of the Erie glacier from the eastern end of the lake.

Lewis (1966) mapped two cross-lake moraines: The Erieau-Cleveland moraine and the Pelee-Lorain moraine. The Erieau-Cleveland moraine extends from Erieau, Ontario, to east of Cleveland, Ohio. The survey for the current study did not extend far enough offshore to encounter this feature, but the trend of the feature toward the U.S. shore (Fig. 21 in Lewis, 1966) indicates that the feature might be related to the broad till platform mapped west of Fairport Harbor. No evidence has been found of an end moraine along the Ohio shore that would tie into this feature. Perhaps the moraine was eroded by streams following glacial retreat or by waves and currents during the early lake stages.

The Pelee-Lorain moraine extends from Pelee Point, Ontario, to Lorain-Vermilion, Ohio. The southern part of this feature was mapped in the survey; however, as with the Erieau-Cleveland moraine, no evidence has been found of an end moraine along the Ohio shore. As an alternate hypothesis perhaps the above named tills were deposited by a glacier limited in extent to the present lake boundaries and thus are younger than the hypothesized tills.

The overall scarcity of glaciolacustrine clay may be the result of Holocene erosion. Erosion of the clay probably took place both subaerially and subaqueously following retreat of the last Wisconsinan glacier from the Lake Erie basin. This hypothesis is supported by modern erosion rates, which indicate appreciably higher erosion rates of glaciolacustrine clay than of till (Carter, 1976). Coarse lag deposits overlying till, particularly in the Fairport Harbor area, and beach deposits on the moraine off Erie, Pennsylvania (Williams and Meisburger, 1982), appear to be good evidence for wave erosion during times of lower lake levels. Lewis (1966) found coarse deposits overlying a smooth till and glaciolacustrine clay surface along the north shore of Lake Erie; he interpreted the surface as a wave-cut terrace. This surface may be correlative with the broad till surface between Fairport Harbor and Cleveland.

Postglacial Sediment.

Holocene sedimentation has contributed to a general filling in and flattening of the lake bottom. The sand deposits at Fairport Harbor, Lorain-Vermilion, and Cedar Point (Williams, et al., 1980) are likely the product of different depositional environments and different ages ranging from the early Holocene, when lake levels were lower than at present, to the late Holocene. Two radiocarbon ages were obtained from wood fragments in core 62 at Fairport Harbor (T.L. Lewis, Cleveland State University, personal communication, 1979). At a depth of 414 centimeters below the lake floor the radiocarbon age is 8250 \pm 145 years B.P.; between 210 and 228 centimeters the radiocarbon age is 4020 \pm 190 years B.P. These ages indicate that the deposit was built up during a rising lake level, possibly the result of a combination of fluvial, deltaic, and beachnearshore environments. Aside from the sand deposits, the finer grained sediments appear, for the most part, to reflect present-day wave and current conditions in the lake, perhaps largely related to fetch and prevailing winds as hypothesized by Thomas, et al. (1976).

V. SUMMARY

The southern Ohio waters of Lake Erie (commonly from 1 to 7 kilometers off-shore) between Conneaut and Marblehead were surveyed in August of 1977 and 1978. Primary survey data consist of 576 kilometers of seismic reflection trackline

profiles, taken in 1977, and 58 vibracores ranging from 0.7 to 6.1 meters long, taken in 1978. About 23 percent of Ohio's part of Lake Erie was covered by the survey.

Shale underlies most of the survey area. The shale is overlain by glacial tills and by postglacial deposits. The surfaces of all the units slope lakeward. In general, the shale is exposed nearest the shore and succeeded offshore by till and postglacial deposits. The shale surface is irregular in comparison with the till and postglacial deposit surfaces and its slopes range from 5 to 20 meters per kilometer. Basal till and flow till were identified in many of the cores. An olive-gray basal till was found over much of the area and a yellowish-gray basal till was found in the Lorain-Vermilion area. Some of the tills contain interlaminated silts and clays (varves?), and three cores contain both basal till and flow till. The tills are made up largely of silt- and clay-size particles composed of quartz and illite. The till surface is much flatter and more uniform than the underlying shale surface and its slopes range from 1 to 4 meters per kilometer. Overall, the till thickens lakeward at about 5 meters per kilometer.

On the basis of cores and echo character on the seismic records, four principal postglacial deposits were defined: sand, muddy sand, sandy mud, and mud. At about 9 kilometers offshore, the stretch from Conneaut to Fairport Harbor is characterized by postglacial sediment thicknesses ranging from 10 to 16 meters; the stretch from Fairport Harbor to Cleveland is characterized by a lack of postglacial sediment. At about 5 kilometers offshore, the area from Cleveland to Lorain is characterized by postglacial sediment thicknesses ranging from 0 to 2 meters; from Lorain to Vermilion the thicknesses range from 4 to 5 meters. In general, the finest sediment is found closer to shore and in shallower water west of Cleveland.

The irregular shale surface in concert with the characteristics of the buried river valleys implies an irregular surface before Pleistocene Glaciation. The relatively smooth till surface is likely the product of several Wisconsinan glacial advances and retreats. The physical variations in the till in the cores, as well as the internal seismic reflectors in the till, indicate that multiple tills were emplaced. The topographic lows on the till surface off Fairport Harbor and Lorain-Vermilion imply stream erosion, and the scarcity of glacio-lacustrine clay implies stream and wave erosion following retreat of the last Wisconsinan glacier. The postglacial deposits accumulated during a rising lake level from fluvial, deltaic, and lacustrine processes. The sedimentary deposits on the lake floor reflect Pleistocene and Holocene processes as well as the configuration and nature of the underlying rock and till. Present-day waves and currents tend to erode and rework the shale, tills, and postglacial deposits in high-energy environments. Subsequent deposition of these sediments in low energy environments is obscuring evidence of previous low water stages.

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APPENDIX A

CORE SEDIMENT DESCRIPTIONS

This appendix contains core sediment descriptions, based on both megascopic and microscopic examinations. Color is based on damp samples as referred to the Munsell color system (Munsell Soil Color Charts, 1954 ed., Munsell Color Co., Inc., Baltimore, Md.); grain size is based on the Wentworth size scale.

The marks on the left side of the stratigraphic log are placed at the midpoints of sampled intervals used for size analysis. The type of analysis is designated by the following codes (size data are tabulated in App. B):

R - Rapid Sand Analyzer (RSA)

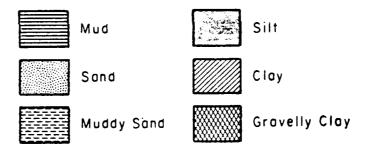
S - Sieve analysis

V - Visual Accumulation Tube (VAT)

P - Pipet analysis

The letter B indicates the bottom of the core.

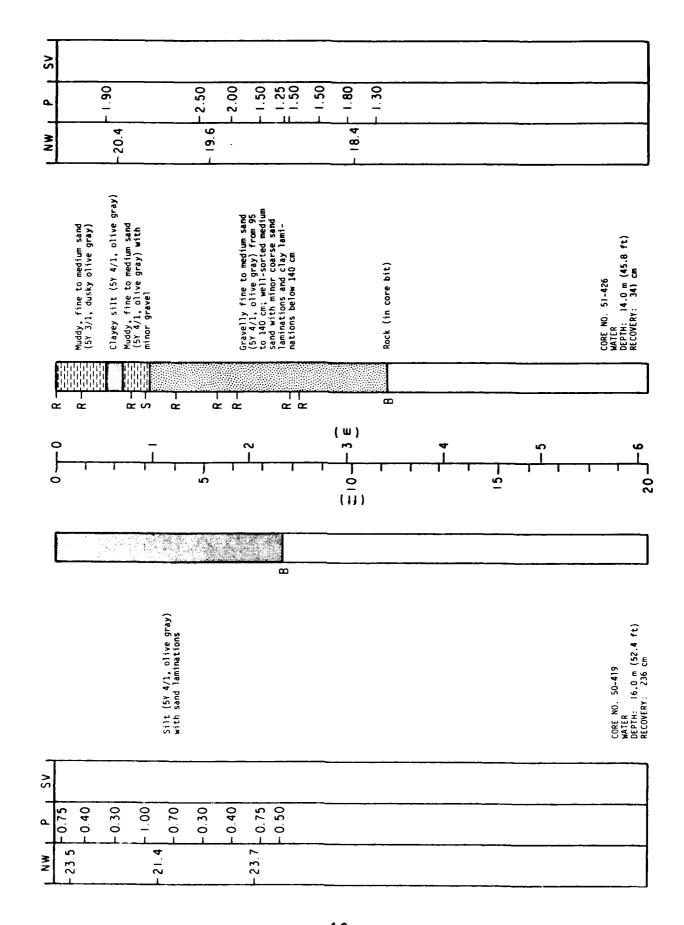
Sediments are grouped into the following six basic categories for logging (minimum unit thickness shown is 20 centimeters):



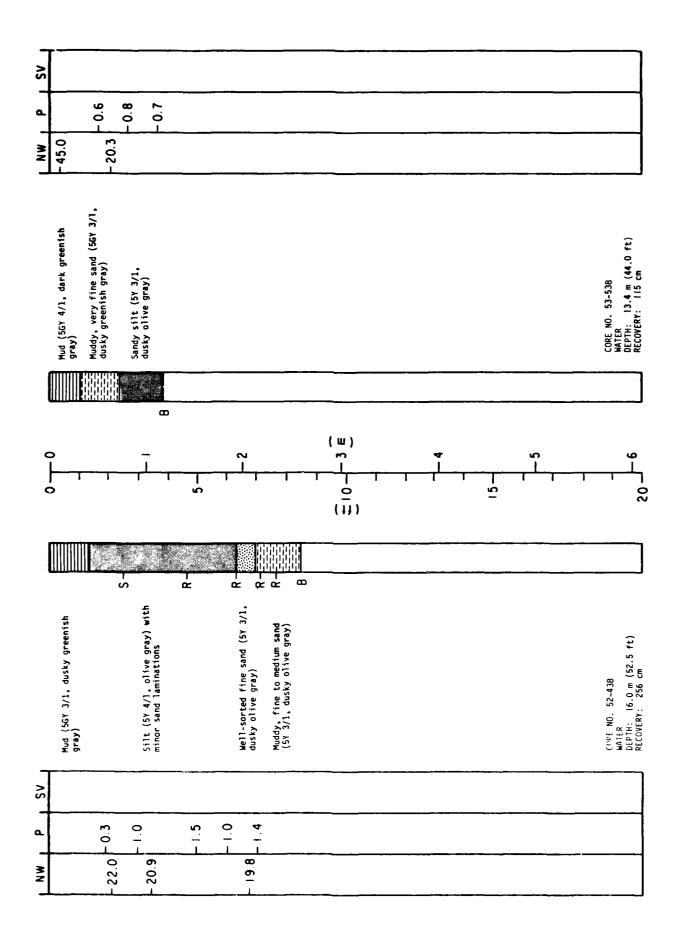
Column descriptions for core sediments are:

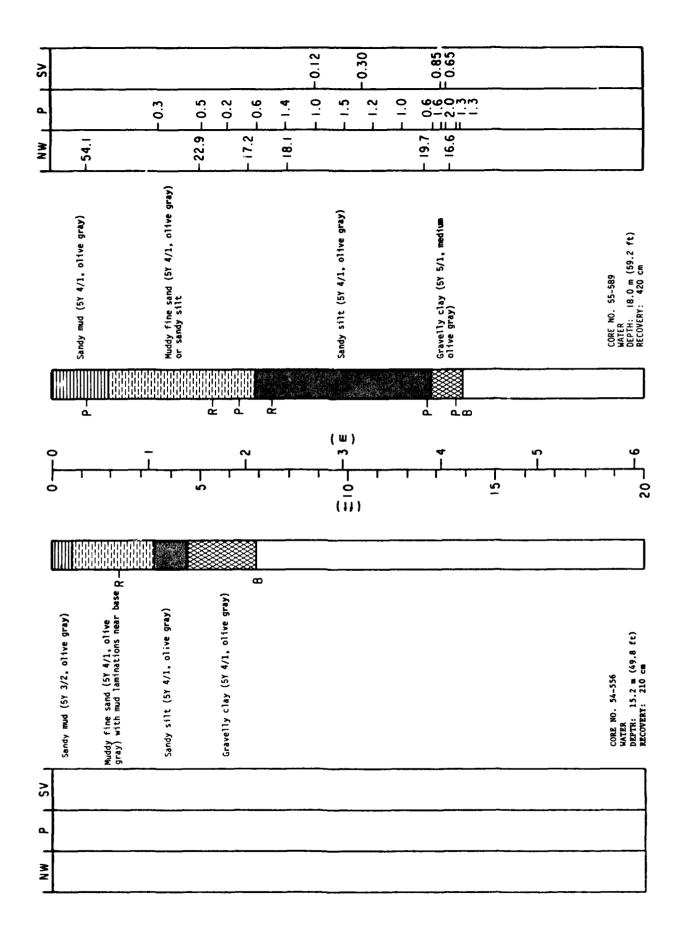
- NW = midpoint of sampled sediment used for natural water determinations (interval usually 10 centimeters long); number is: weight of wet sediment minus weight of dry sediment divided by weight of wet sediment times 100.
- P = penetrometer measurement (in tons per square inch or kilograms per square centimeter).
- SV = shear vane measurement (in tons per square inch or kilograms per square centimeter); PF = plastic flow.

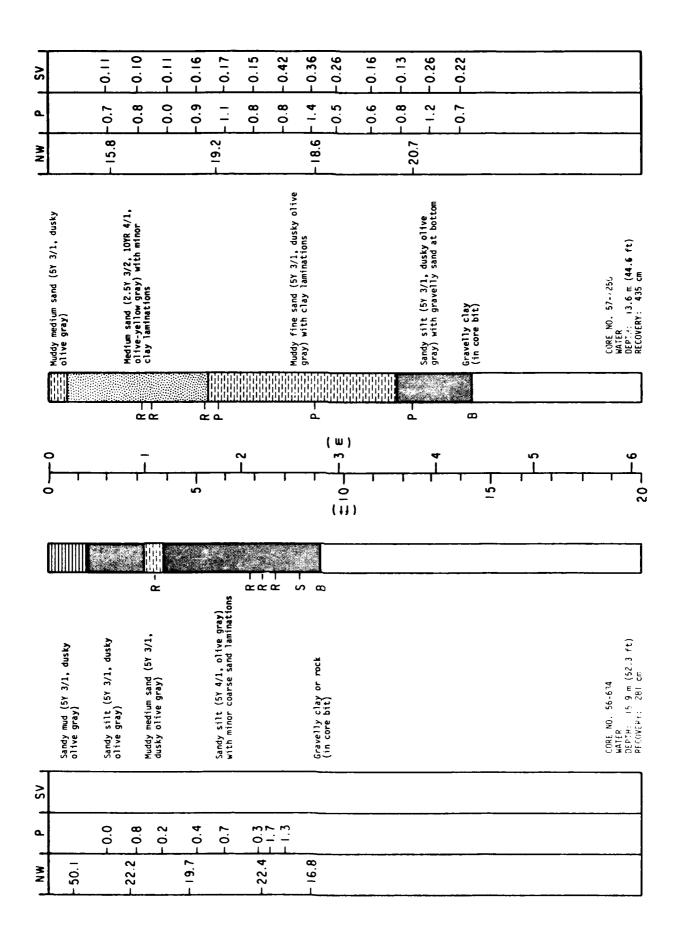
Water depths are the surveyed water depths; these depths were about 1 meter above low water datum (LWD) for Lake Erie.

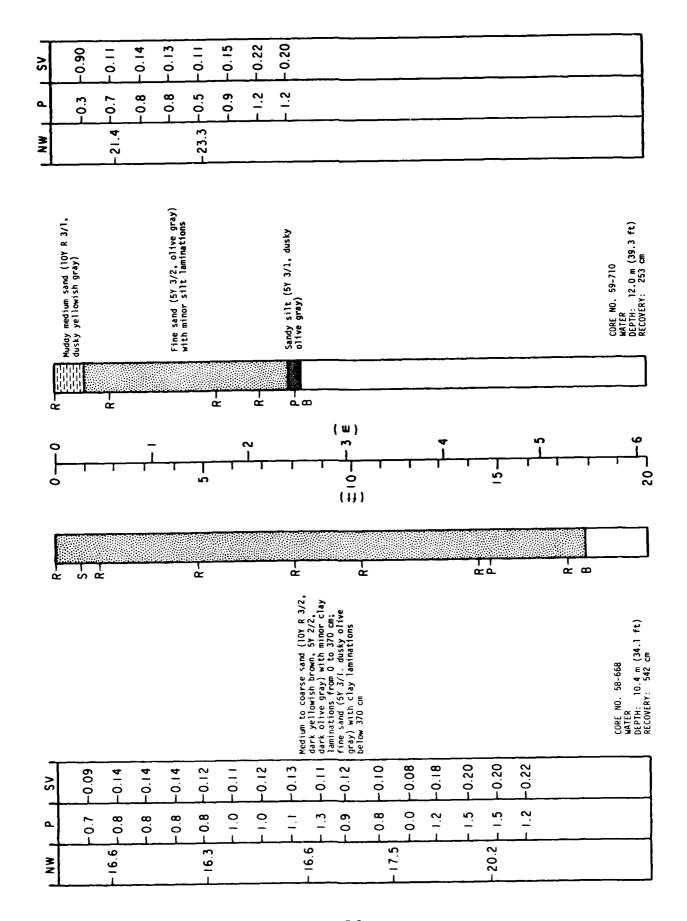


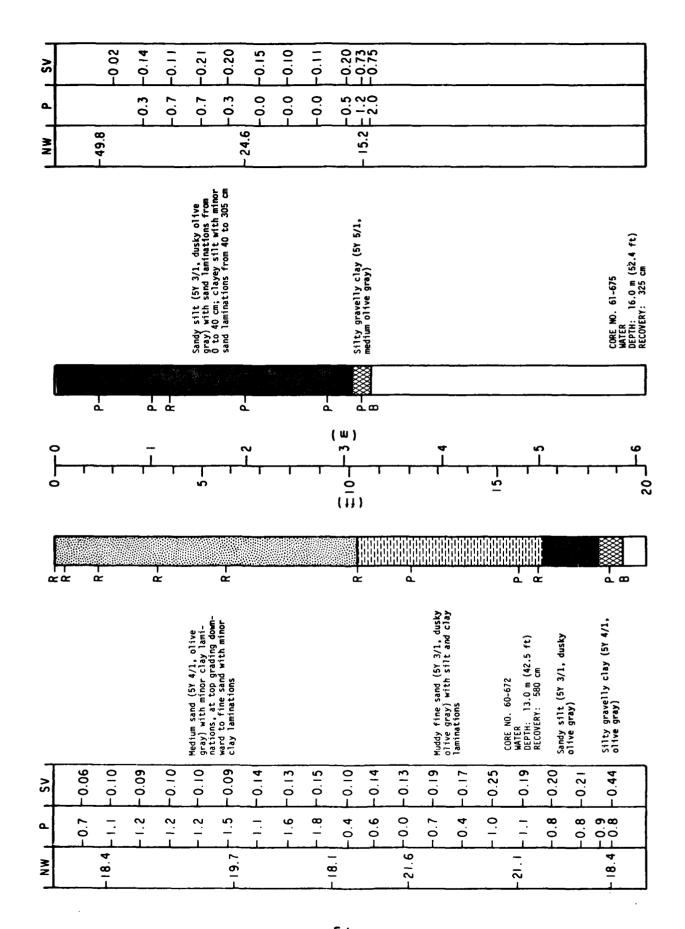
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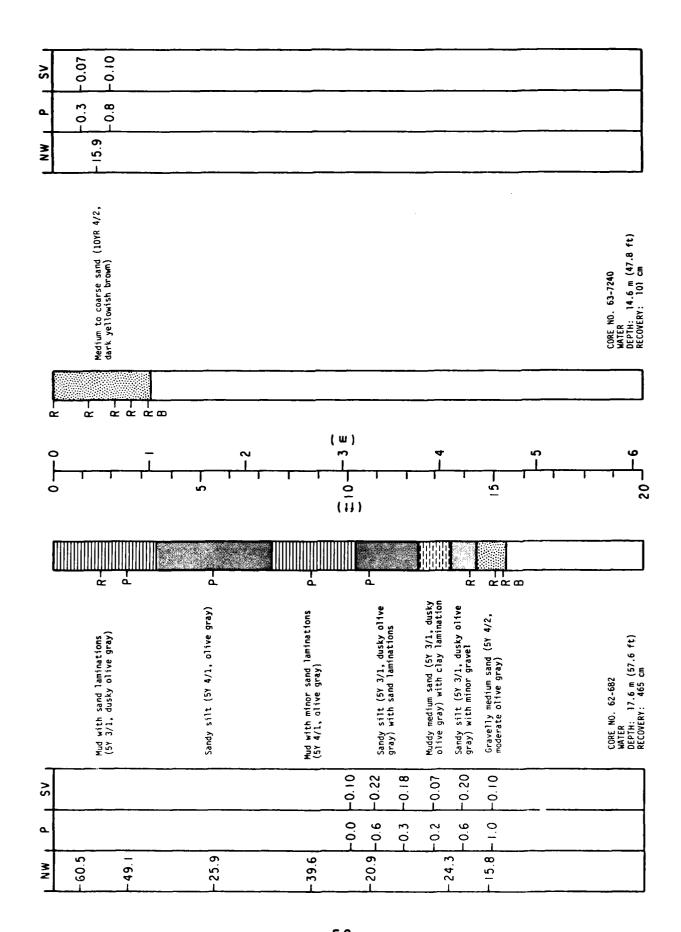




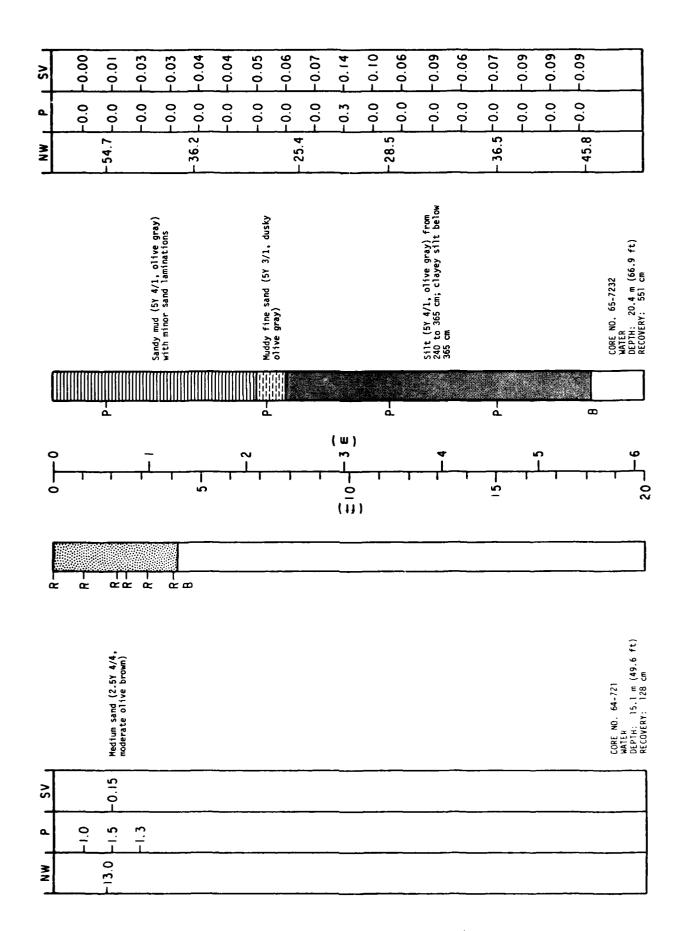


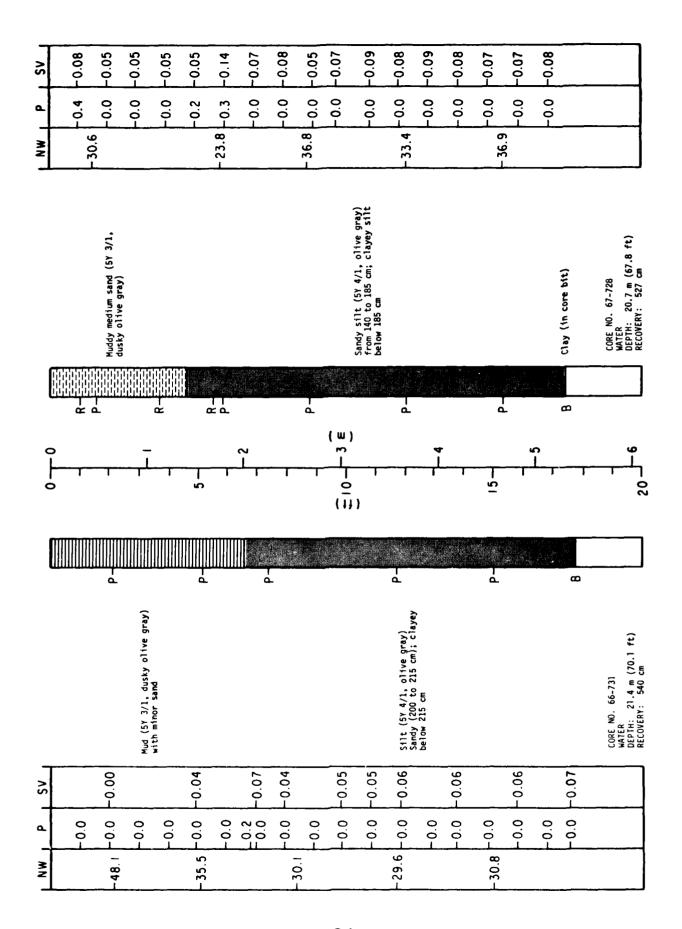


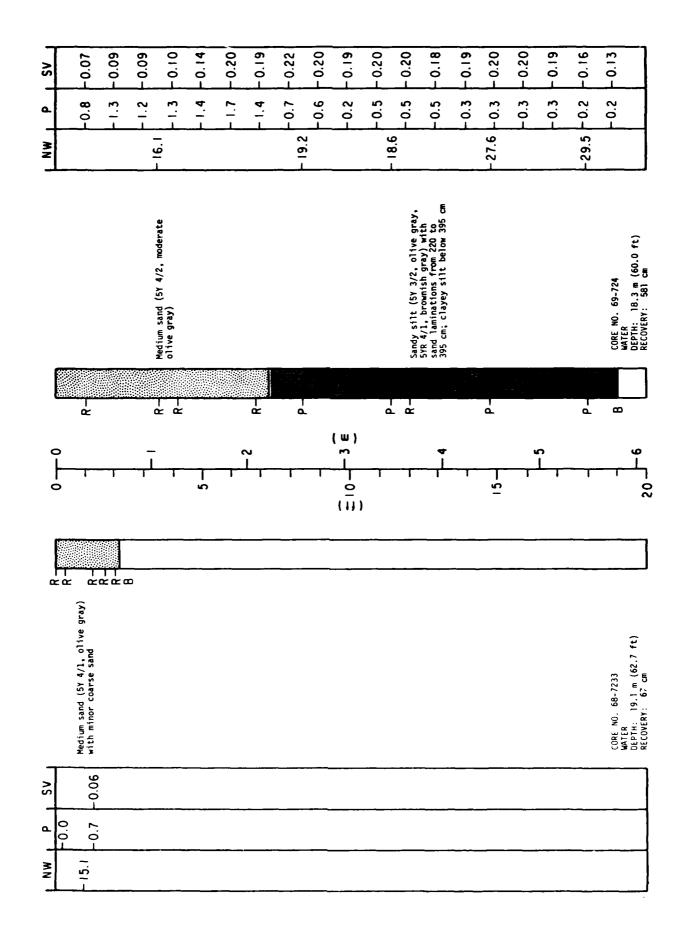


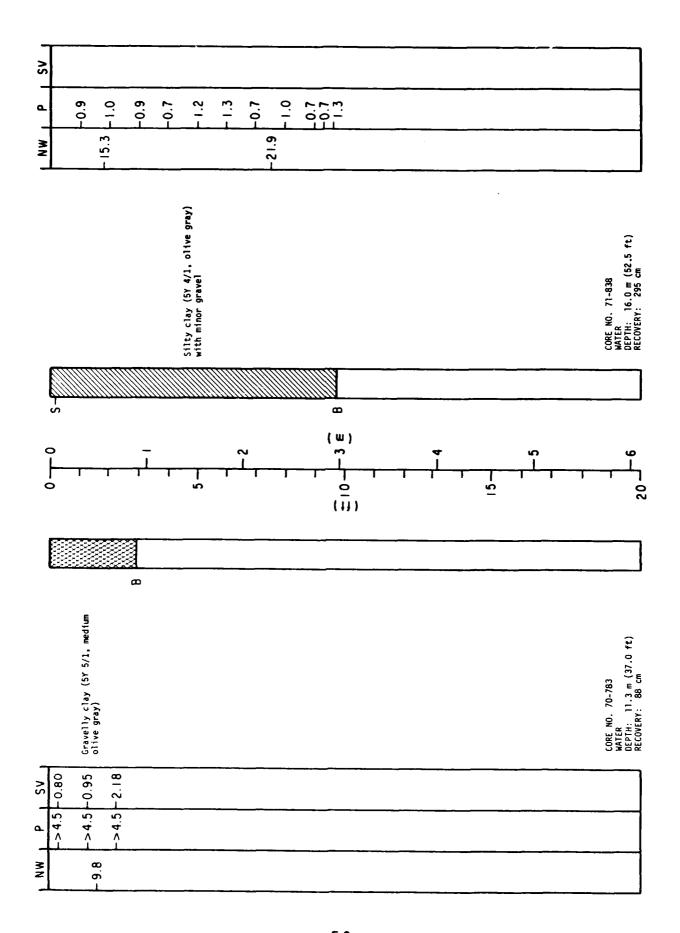


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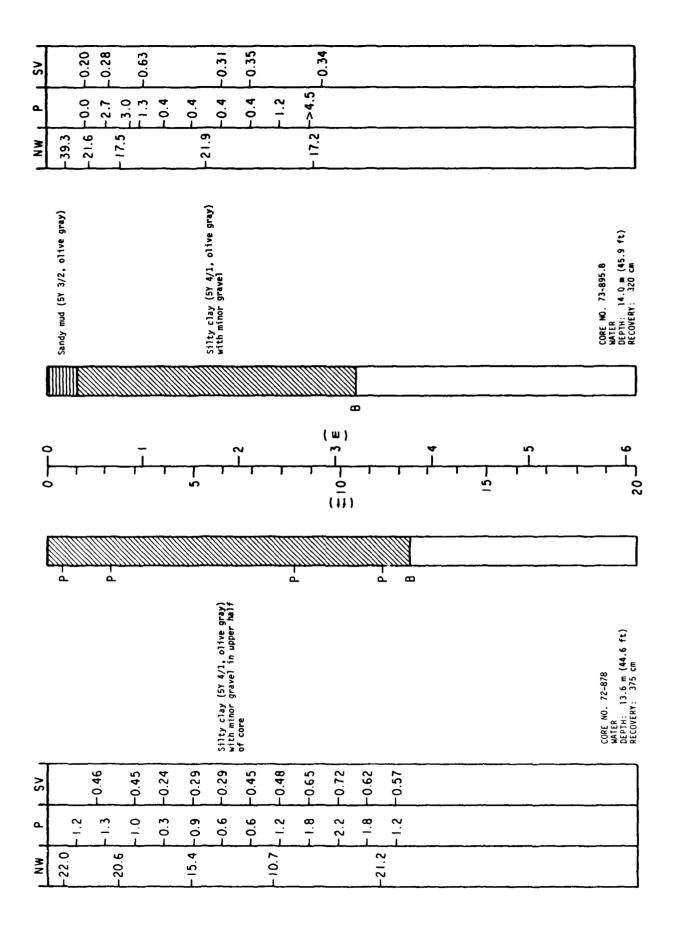


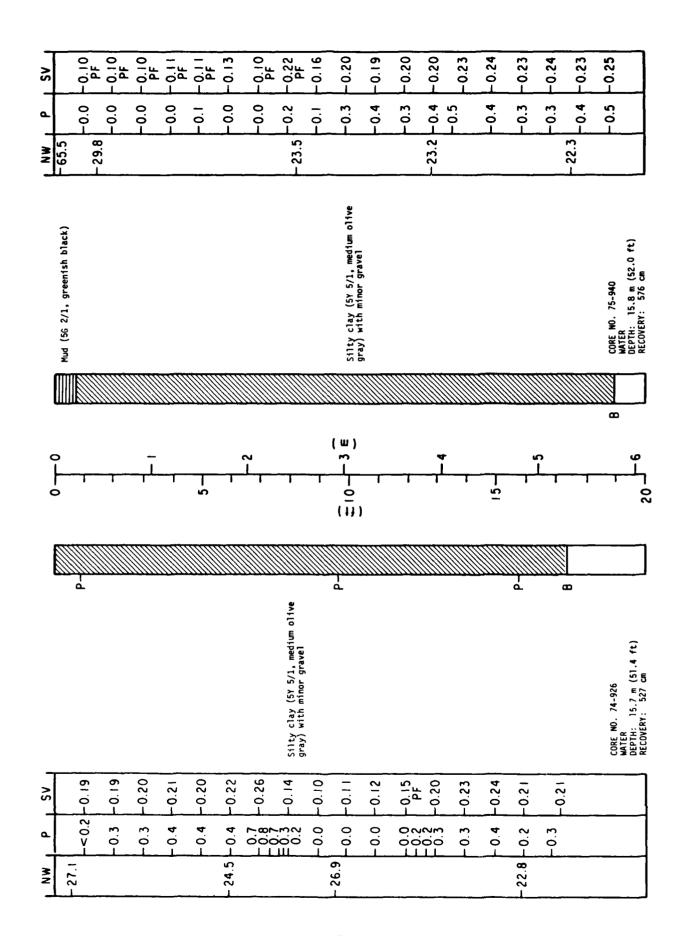




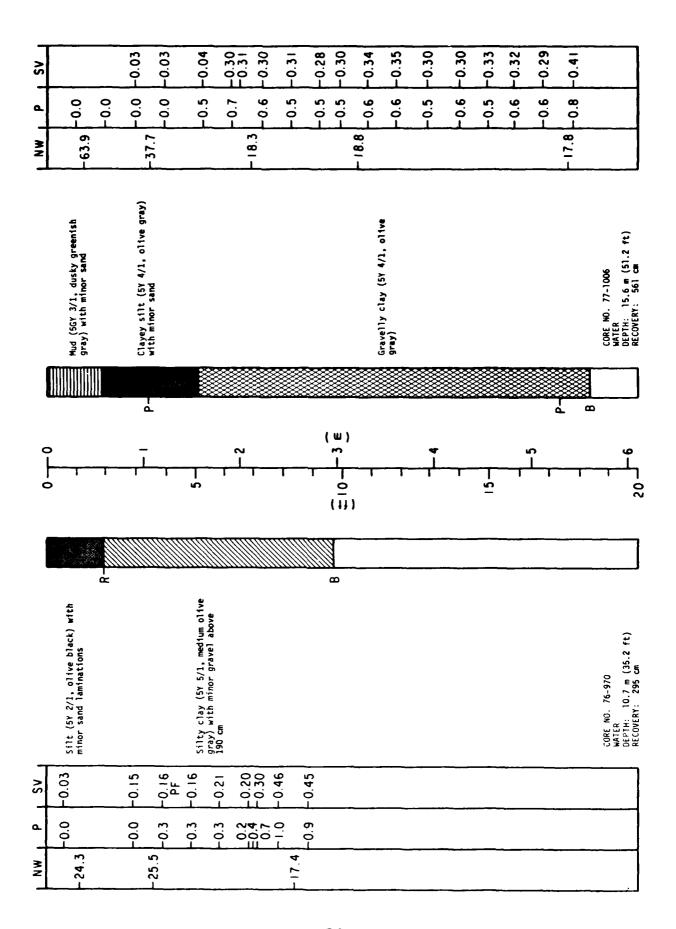


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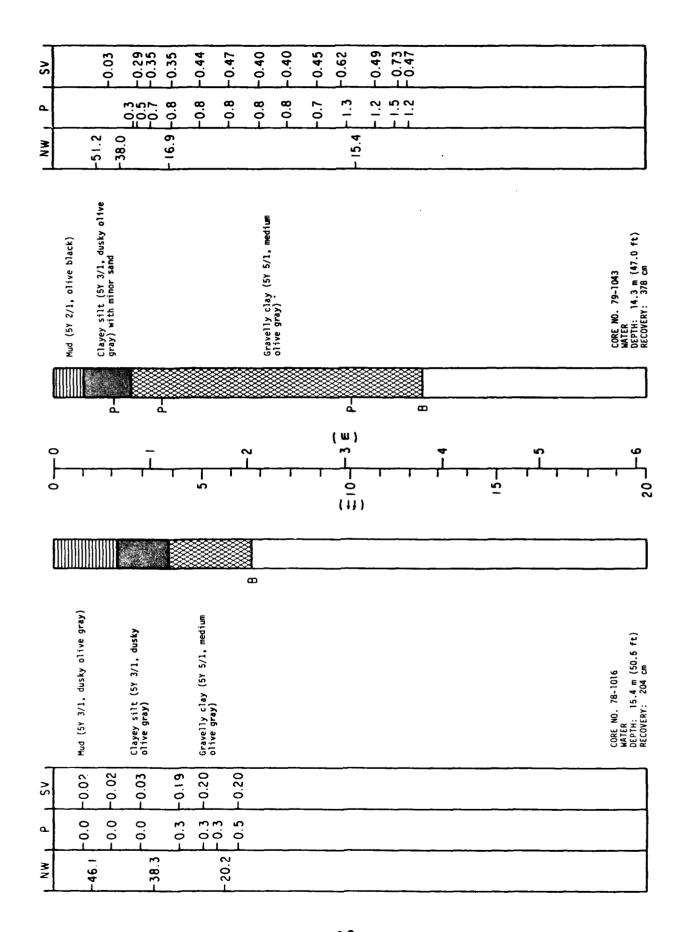


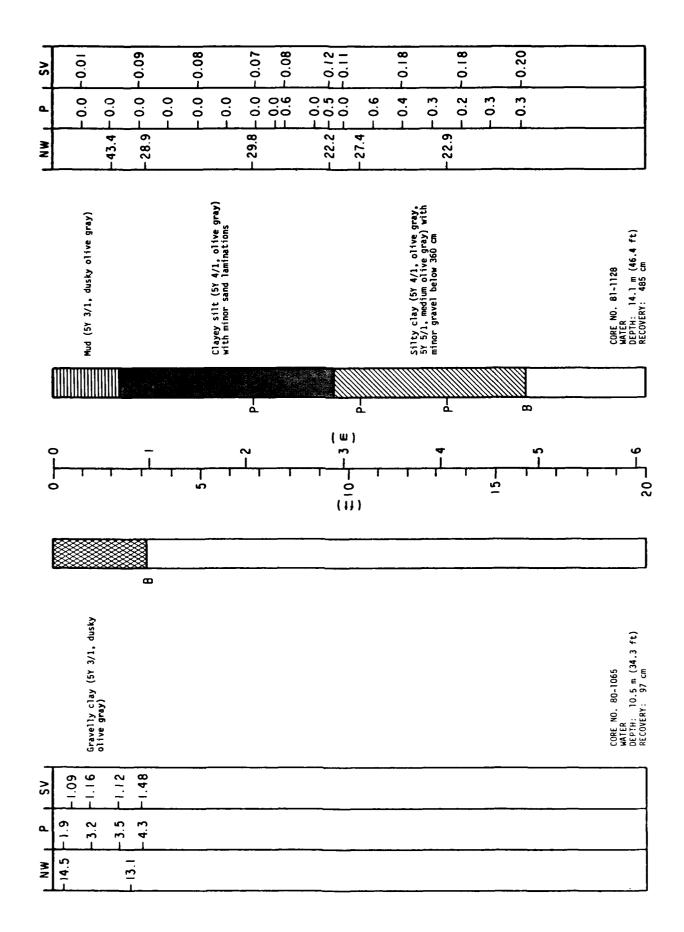


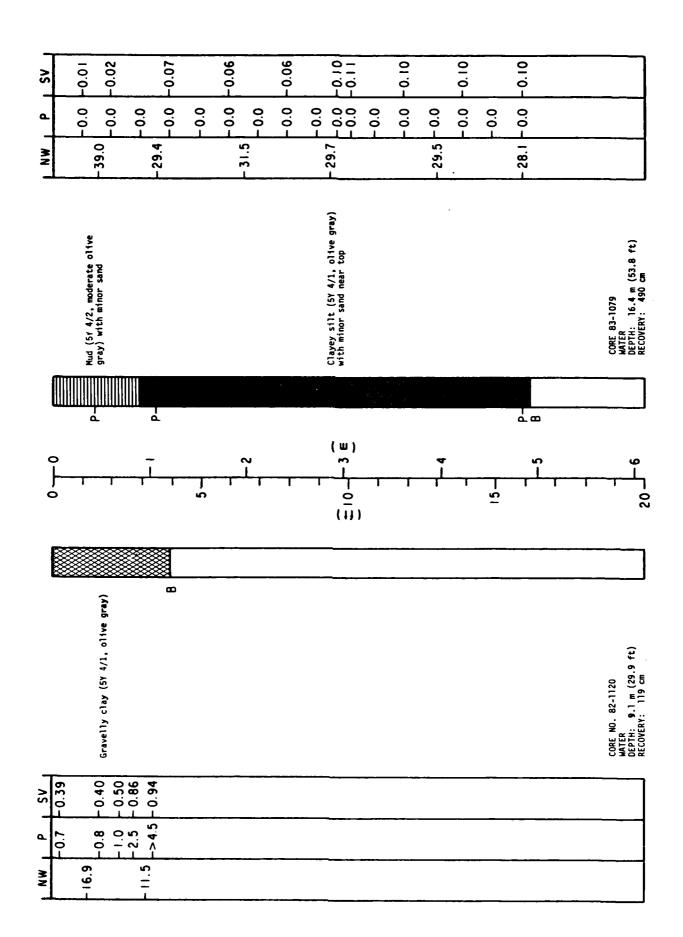
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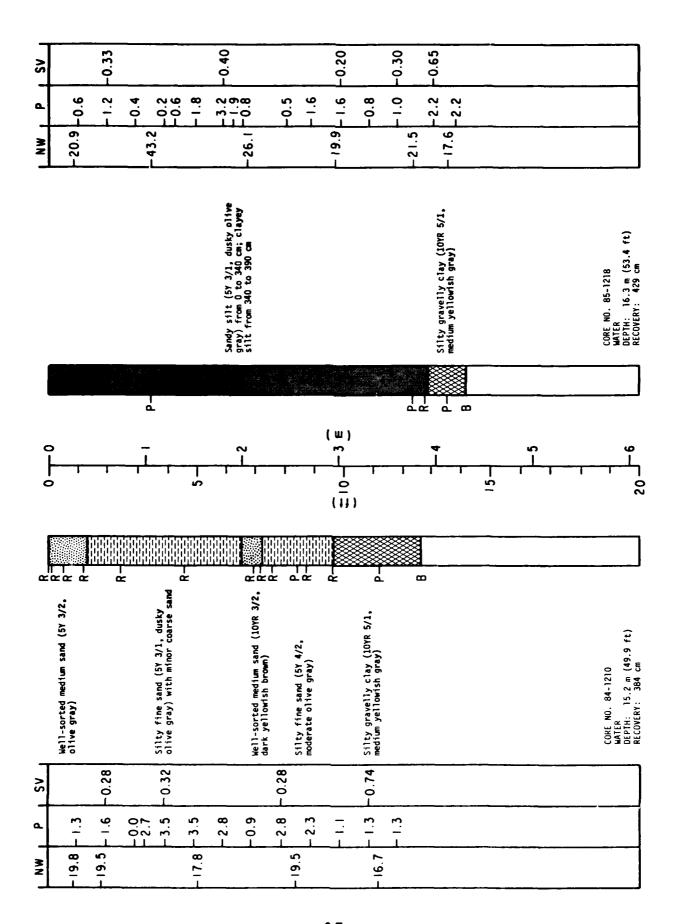


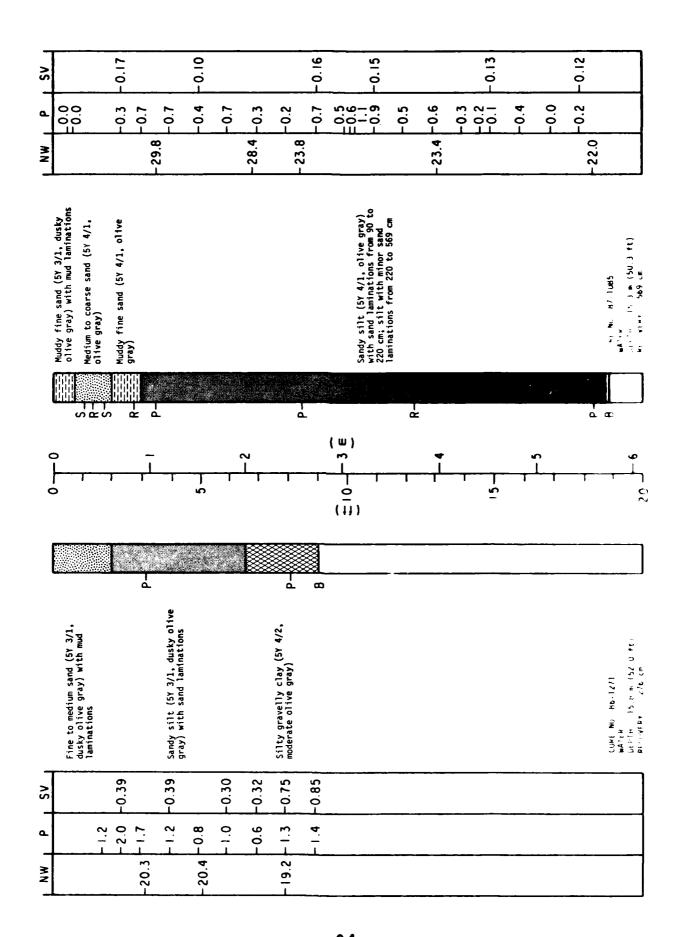
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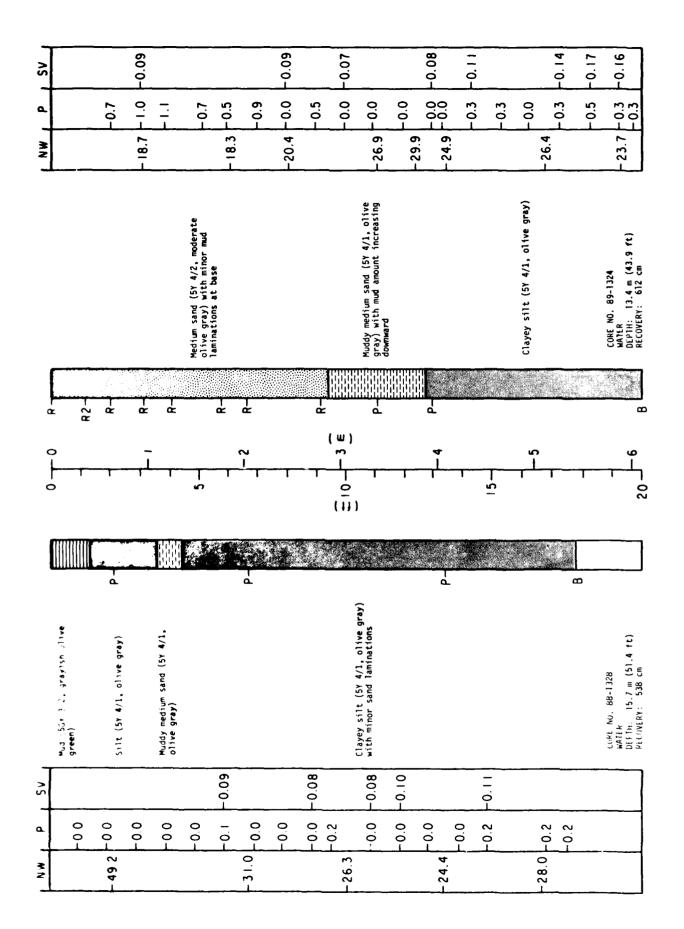


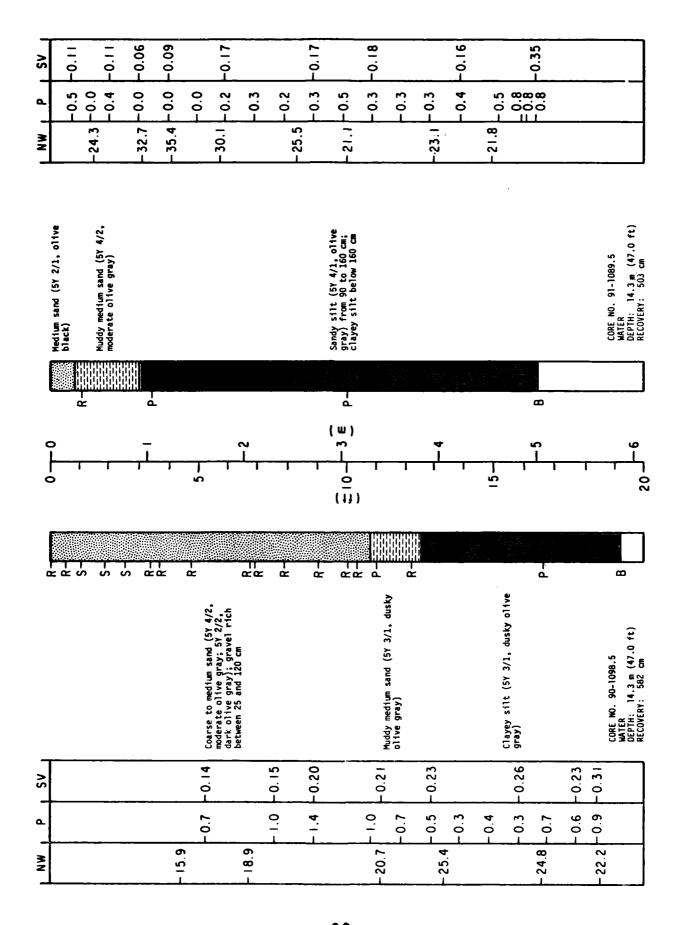




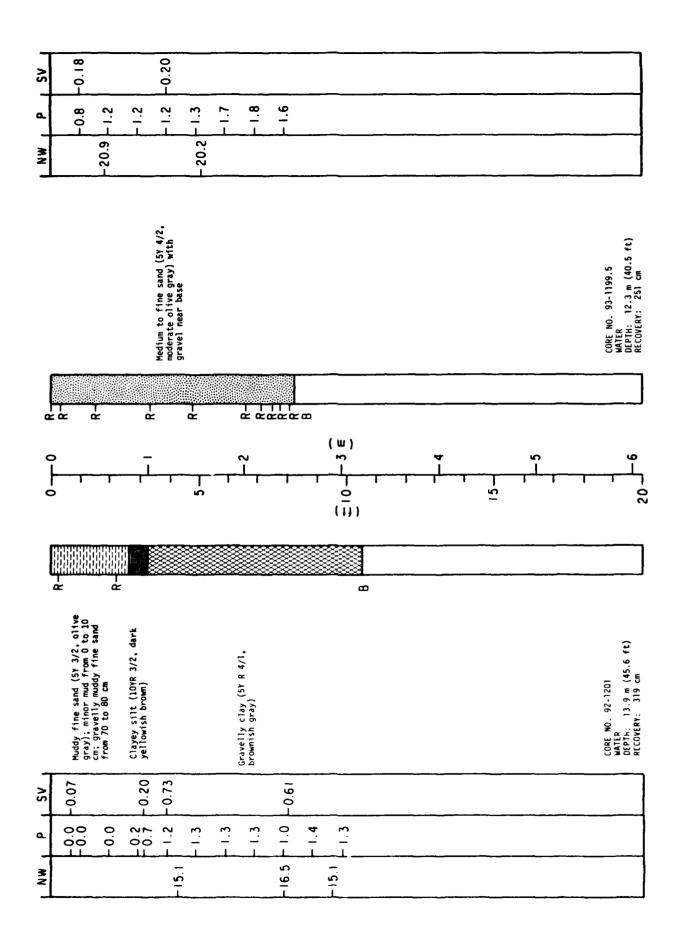


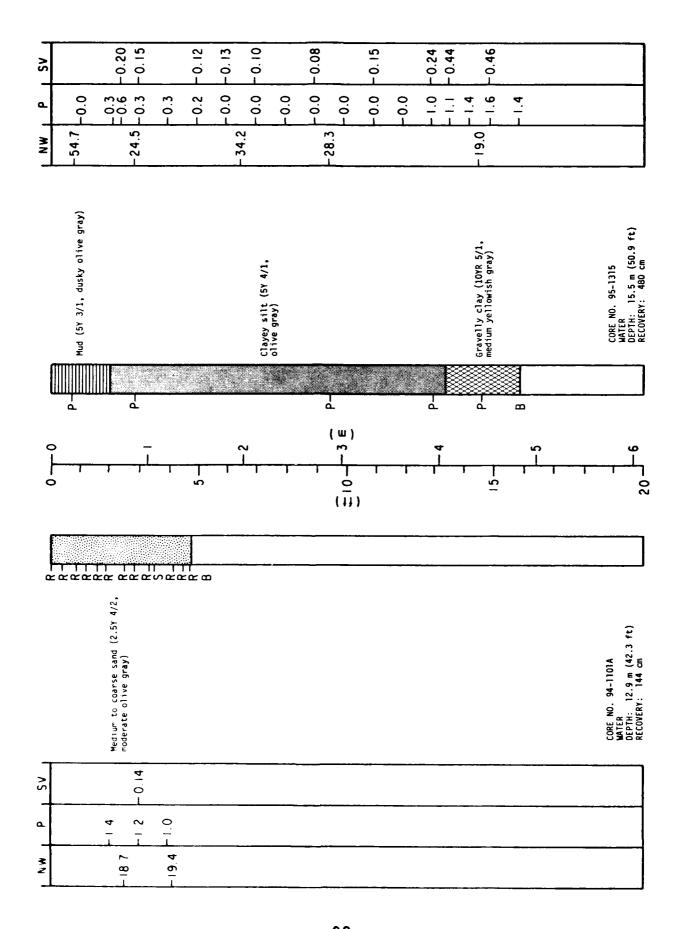


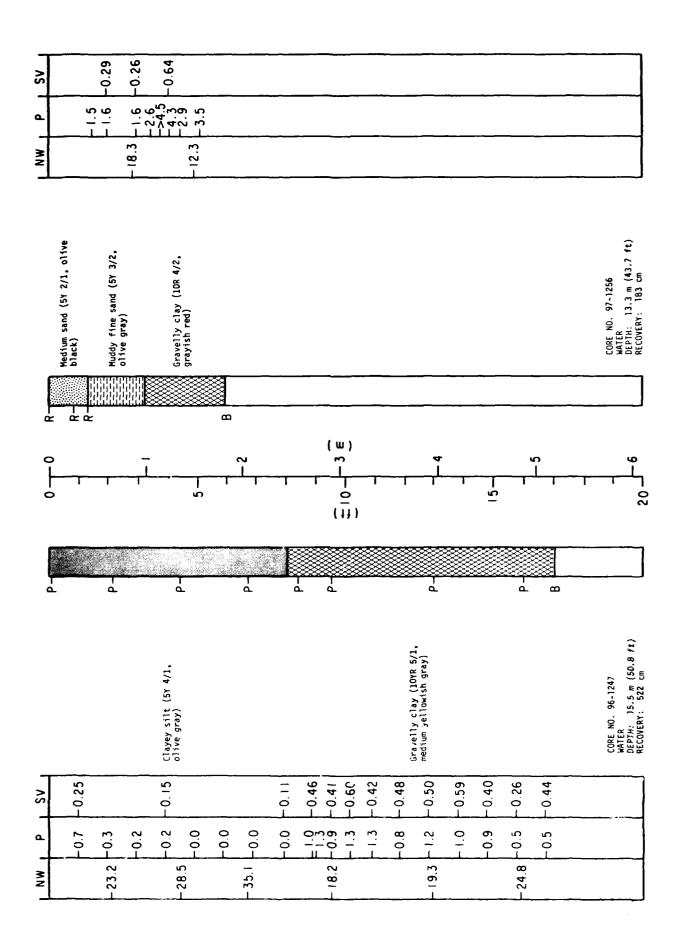


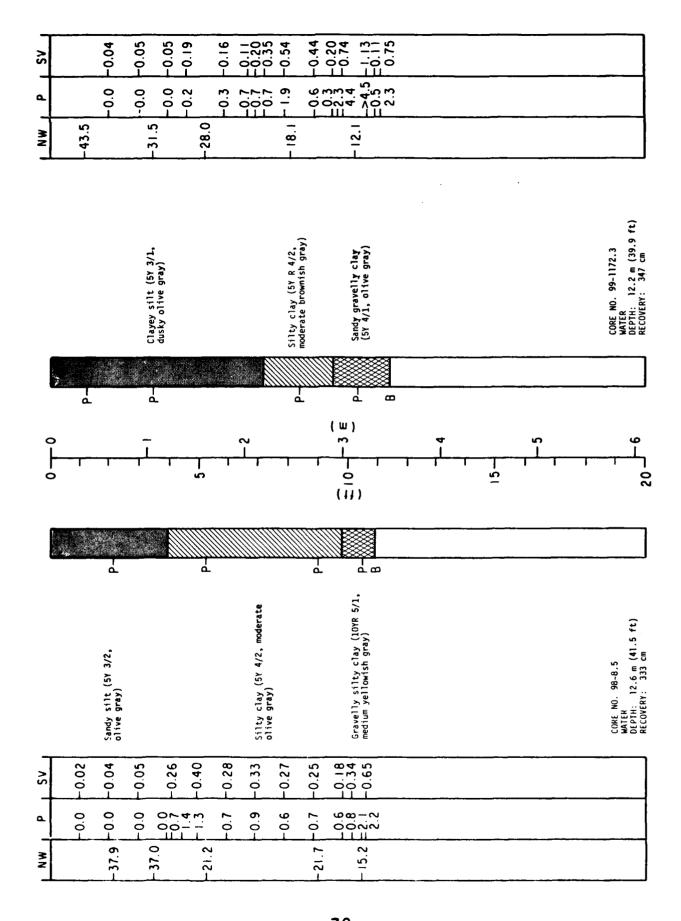


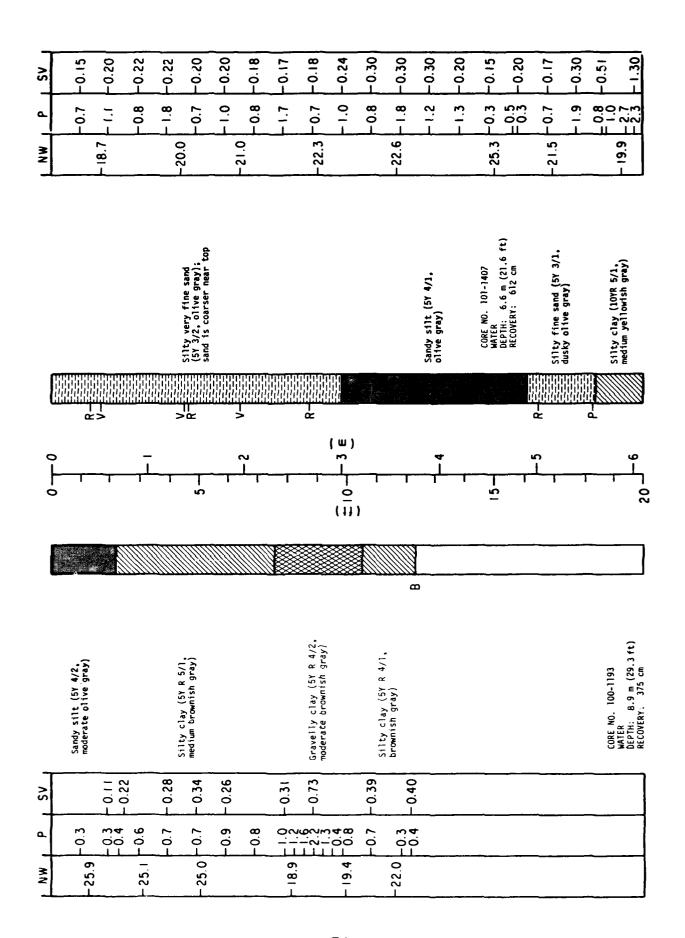
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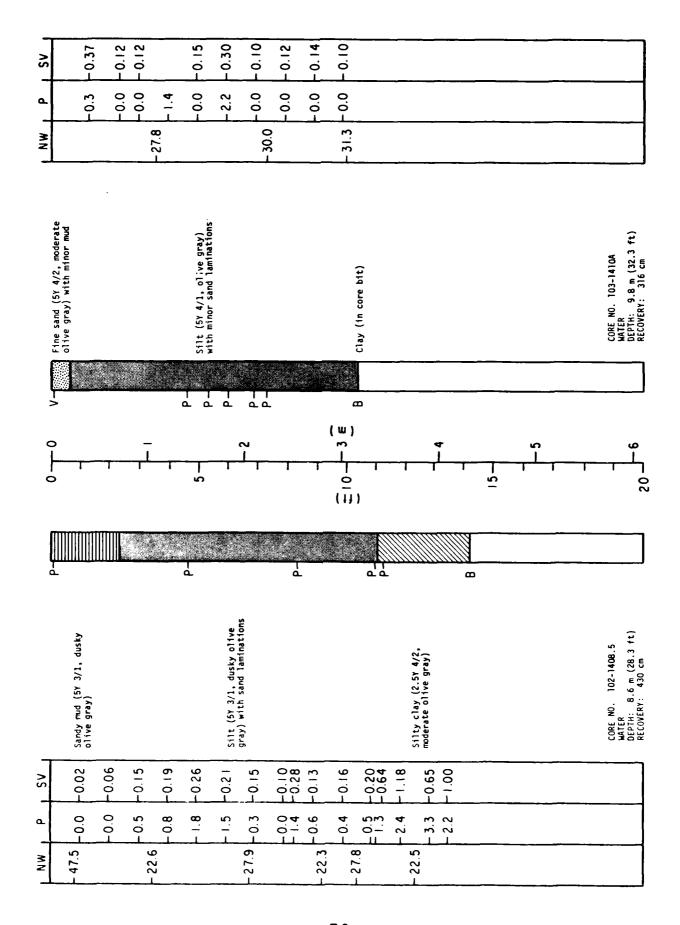




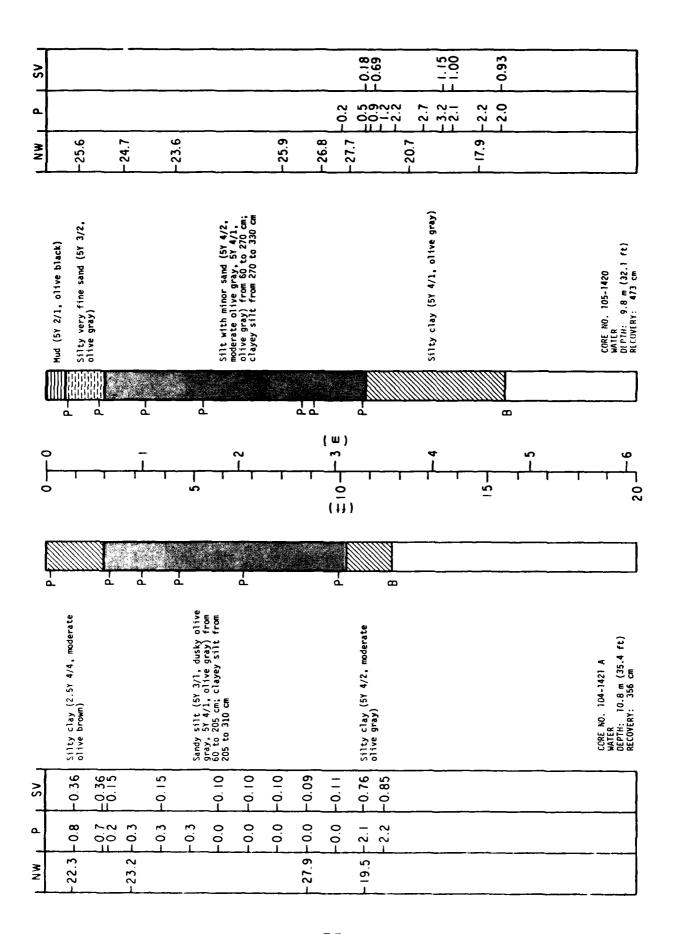


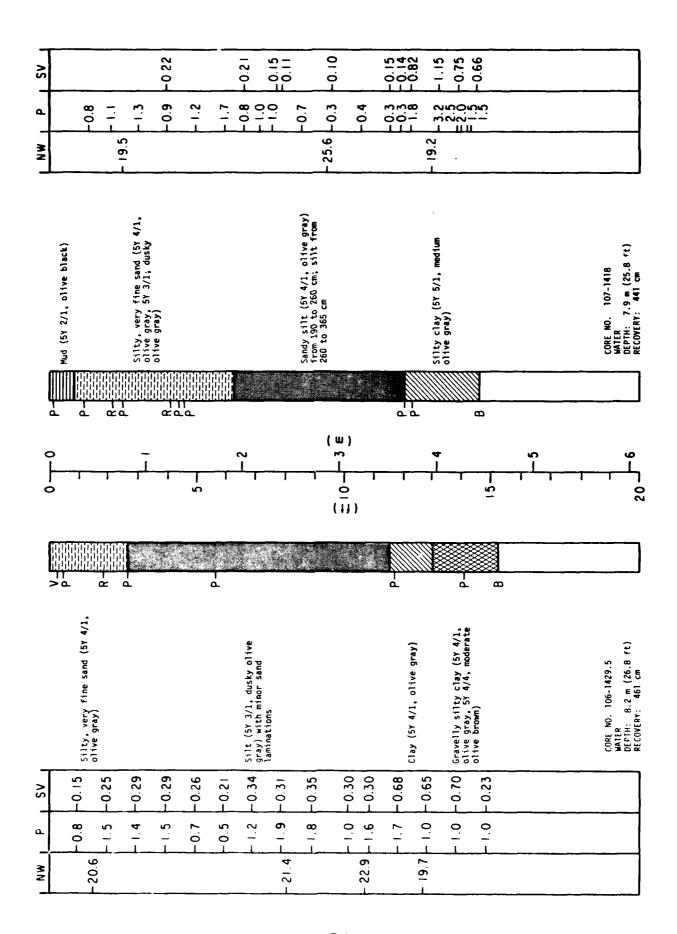






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APPENDIX B

GRANULOMETRIC DATA AND CUMULATIVE CURVE PLOTS

The samples in this appendix are identified by core number and sample interval below the top of the core. Locations of the samples in each core are shown in Appendix A.

1. Rapid Sand Analyzer (RSA).

Data include the frequency and cumulative percent at 0.5-phi intervals. Also included are median, mean, standard deviation, skewness, and kurtosis for each sample. Experience has shown that grain-size values from RSA analyses are consistent and slightly coarser than results of sieve analyses of identical samples; therefore, empirical relations for converting RSA means and standard deviation to sieve analyses equivalents have been determined. The relation-ships, developed from RSA and sieve analyses at a 0.25-phi interval are (Williams, Prins, and Meisburger, 1979):

Mean:
$$\overline{X}_{\emptyset}$$
 sieve = 1.0735 $\overline{X}_{\emptyset RSA}$ + 0.1876

Standard deviation:
$$0_{\emptyset}$$
 sieve = 1.4535 0_{\emptyset} RSA - 0.146

2. Sieve Data.

Data include frequency and cumulative percent at 0.5-phi intervals from samples estimated to have gravel percentages over 10 percent.

3. Cumulative Curves.

Sieve data plotted at 0.5-phi intervals.

4. Pipet Analysis.

Percentages of sand, silt, and clay are included for each sample.

5. Visual Accumulation Tube (VAT) (Guy, 1969).

Percentages of sand and silt-clay as well as mean grain size are included for each sample.

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		DISTRIBUT	-	ON OF	SIZE		CLASSES 1	BY FR	FREQUENCY		PERCENTAGES	AGES		(RSA)Continued	ntinu	eq		
2601 109E	FERENCE 1.000	005.	•	500	3126	. 50	CLA59	c	1,000	•	•	_	-	# 0 1 4	Ξ	MEAN	CAL PA	RAME TERS
ç	. ((4) 2.000	1.4.1	1.000	707	• •		• •	.177	.125	0.0	203	2	ī	1	Ĭ.	H.	i.	
0	274-275 0.00	9.0	00.0				31.70	37.80	13.20	Š			•	2.19	.219	2.37	.194	14.
	315-30, 0.00		6	€.		1.2	•	~	5:	7.62	•	Č		2.21	.216	2,28	.20e	.77
	310-3200.00	•	000	0.0	ç	•	17.26	28.87	-	::	×	ō		•	.183	2.53	.173	57.
	364-37/ 0.00	•	10.70	24.84	•	13.13		•	1.11	5.12	•	ē		.78	.584	1:1	797.	1.27
	32-33 0.10			. 3	-	7.40	•	٠	è	5.54	•	ō		2.19	•12.	2.23	.213	.72
	0-10 0.00			0:-	-	2.09	16.67	43.12	33.66	0000	0.0	ō		7	200	2.24	.211	.54
	04-76 0.00	•	0.00	00.0	۰	1.20	•	40.5	30.10	4.56	8.55	ō		2.57	.168	7.04	.160	70.
	100 0.00	·	0.00	00.6	ŷ.		20.55	32.69	21.12		~			7	961	٦.		.05
	10-11 0.00	E	0.0	٠.	-	2.65	21.65	43.83	24.18	4.12	0			2.27	.207	7.27	.207	57.
	80-50 0.00	_	2.50	===	. 36		22.99	39.97	34.08	•	0.00	000		~	.207	-	155.	•
	110-104 0.00	_	4.23	1 . A S	3.16	2:53		31.74	27.23	3.01	0.0	0		.23	.213	2.05	502.	70.
	1-1-150000	•	1.90		0.0	Ξ	: 0.32	\$6.17	20.03	11.74	٠.	00.0		.38	261.	₹.	191	. 0.7
	200-5010.00	•	6.	~9.		9.16	~	25.65	•			•		. 35	661	~	197	.75
	210-2200.00			1.30	2.30	5.13	•	29.90	34.10	-	2.61	Ö		2.07	190	7	193	1
		C	00.0	.0	•	1.34	17.14	25.51	٠	. =		Ö			169	2.55	17.	
		•	=	5.7	0	10.88	13.79	10.45	10.64	14.17	10.5	ē		1.79	289	•	. 311	
		•		~	2.28	5.4		20.53	27.74	74.6	•	ē		2.39	161	2.38	761	6
	100 0.00	•	3			16.24		29.37	-	•	000	0.0		•	267	76.	. 261	100
		•	ē	00.0	-	20.27	2.0	22.93	A. 56	9.66	¢	Ē		19.	280	1.90	.267	3
	20-30 0.00	•	v.	c •	2	27.45	23.48	17.74	9	6.83		Ē		79.1	321	1.70	308	54.
	34-15 0.00	Ξ	-	•	-		~	9.53	A. 17	2.81	9	٩		. 7	5.85	80	260	
	00.00 40.00	` ^		24.61	200	16.18	12.82	7	7.15	3.01	9	Ö		•	526	90.1	471	30
	50-60 0.00			16.05	7.5	17.54	2.0	63.0	7.01	000	Ö	000		1.19	917	1.30	407	90.
	70-90 0.00	E	ē	. A.	1.77	21.00	3	21.28	9.03	000	•	•		1.72	305	8	285	9
	80-49 0.00	•	c	٠.	٠.	29.08	15.46	20.23	9.45	5.69	00.0	00.0		1.7	300	1.85	.277	.56
	100-102 0.00	~	Ē	16.33	٥. ٧		13.91	15.21	9.86	3.73		Ē		1,35	.393	1.17	£ 77	01.0
	120-130 0.00		~	٩.	~, -	4.67	29.00	25.29	-	8.61	7.34			2,25	.210	2.21	.216	Ξ.
	130-140 0.00	•	Ġ	1.53		Ÿ.	17.24	30.13	17.17	40.0	٠.	0.00		2.25	. Z.	2.26	.208	.75
	143 0.00	•	Ē	0.00	.55	2.43	24.10	31.69	3	67.6	•	0.00		2,37	. 76	2.45	.185	74.
	10P 0.00	•	٠.	11.38	17.55	15.41	13.40	15.05	13.39	•	•	000		1,45	. 365	67.	157	0
	24-19 0.00		9.08	10.0	3	0.20	3.	19.00	~	9.0	8.27	ē		2.01	977	1.86	.272	1.13
	00.0 19-07		-	5.05	4.1	3.46	14.32	16.52	. 25	10.01		ě		2,39	101.	٩.	. 235	200
	3>045 000	ح	0.0	0.63	٠.	010	٥, ٠	3.5	3	30.86	00.07	000		3.08	911	•	124	.50
	140-1500.00	٠		2.42	1.37	2.26	5.20	9.77	1.14	45.24	N	00.0		3.10		2.40	136	.75
	200-1700.00	•	ē		1.42	2.07	Ť.	5.68	09.0	37.43	30.65	0.00		~		0	-11	.65
	\$c1-533 0.00	-		1.42	. 0.	7.54	~	10:10	.59	~	•	0.00		2,93	.132	2.75	91.	~
•	\$0.00 0.00	•	3.	. 47	٠٠	5	2.19	2.0	0	39.73	30.52	00.0		N		=	==	.50
•	00.0 0.00	•	0000	0.00	•	0.00	۲.	11.96	~	2	•	0000		3,25		3.16	.112	.51
_	120-130 0.00		•	٥.	1.52	. 1.	2.1.	1.0	27.10	•	Ť.	0000		Ñ	• 100	3.09	911.	9

	Date May 1979		# Date	Cumulative.	Per Cent Passing																										+			
				16705	Cumulative Per Cent						2.38		4.35	6.17	18.8		15.53	27 98	22:37	31.62	63.62		53.32	67.73		76.91	81.14		81.60	81.92		84.83	93.44	100.00
מנייונייו עימויוזים	d by_		YSTS OF SAND	Retained on Sieves	Per Cent						2.38		1.97	1.82	3,64		6.72	3.4.6		10.63	9.80		9.90	14.40		9.19	4.23		3	0.32		2.92	8.60	6.56
מנייני שנייניים	Collected by		SIEVE AMALYSIS OF Analyzed by		Grams						1,63	***	1.35	1,25			4.61	-	3.11	7.29	6.72		6.79	9.88		6,30	2.90		0.3	0.22		7.8	5.90	4.50
	KIE - 0HIO	76-77 CB	95	U.S.	Mesh	1 3/64	7/8	3/4	8/5	1/2	9/8	5/16	1/1	3 1/2	.,	9	,	80 9	12	71	1 16	20	25	2 2	0.7	5 45	9	70	200	120	1.0	170	\$30	Pan
	CONS LAKE ERIE	ORE 52 76	Weight of Sample 70.05	Screen	Opening NY	26.670	22.400	19.200	16.000	13.550	9 520	7.930	6.350	5.613	3.953	3, 360	2.794	2.362	1 700	1.400	1.180	850	017.	909	.425	. 355	. 250	.212	.180	.125	. 106	86.	.063	00.00
٠	Project ICONS	Recaris CORE 52	Veight of		•	-6.67	-5.60	-4.80	60.7-	-3.68	200	-3.00	-2.65	-2.50	-2.25	27:1-	-1.50	-1.25	-1.00	-0.50	-0.23	+6.25	+0.30	£ 2	+1.25	+1.50	+2.30	+2.25	+2.59	+3.00	+3.25	+3.50	03.7	
B	Date May 1979		Date	Cumulative	Per Cent Passing																										<u>-</u>			
			,	Sleves	Cumulacive Per Cent				-A	7.87	8.13	8,72	80.6	9.41	70 01	nava.	11.93	12 70	14.10	13.41	17 71	42454	16.58	21.75		30.52	38,63	50 67	93,00	53.70	40 00	79.88	88,15	100,00
CERC SPULIENT ANALYSIS	by		SIS OF SAND by	Retained on S	1 1					5.87	1.09	0.59	0.36	0.33	37 1		1.07	0 22	0.,,	0,71	90	7.7.4	2.17	5.17		8.77	8.11		1514	10.70	: ;;	26.19	8,26	11,85
	Collected		SIEVE ANALYSIS OF Analyzed by		Grams					19.82	3.70	2.00	1.20	1.12	00 4		3.60	6	700.7	2.39	07 €	71.7	7.32	17.43		29.59	27.35		14:72	36.10	36 00	88.35	27.88	39.99
CERC SFL	ا	- (1	- Z		I					,		5/16	•	3 1/2	.7	9	1	80 (2	12	14	18	20	25	355	07	\$ 65	60	70	100	120	170	200	230	Pan
CERC SFI	- OKIO	RIE, PA 0 - 105 cm	<u>i</u>	5 1	Mesh	1 3/64	7/8	3/4	R/S	1//2	3/8	5/2	1/;				Ц	1		Ш	╧			1					\int		1			
CERC SFI	E ERIE - OKIO	Locationy Staple No. ERIE, PA Remarks CORE 51 80 - 105 cm	0.41 8r.	Screen	**	26.670 1 3/64	H			+	3/5 0 3/8	+			4.760	3,360	2.794	2.362	1.700	1.400	1.130	.850	.710	. 500	.425	355	. 250	.212			1	.075		0

	Date May 1979	s Date	Cumulative	Per Cent Passing												1																	
			eves	Cumulacive Per Cent							0.94	2.09	3.20	15. 53		26.74	35.06		40.52	44.21	48.71	59.22		82.31	24,67		27.182	99.10	79.66	12:22	99. B4	Tro var	
CERC SEPTIENT ANALYSIS	l by	SIS OF SAND	Retained on Sieves	Per Cent							0.94	1.15	1.10	10.0		14.54	8.31		5.46	3.69	4.50	10.51		23.09	12,36		2.72	1.48	0.57		77.0	VAVO	
ERC SPULLE	Collected by	SIEVE AMALYSIS Analyzed by_		Grans							1.21	1.48	1.42	3		18.70	10.70		7.02	4.75	5.79	11.53		29.70	15.90		3.79	1.90	0.73	,	22.0	128 63	-1.36
O,	1.AKE ERIE - OHIO 0. 25 cm	<u>i.</u>	U.S.	Mesh	1 3/64	7/8	3/4	578	1/2	8/2	5/16	1/4	3 1/2	,	9	~ 0	10	12	14	18	25	8	7.0	45	60	70	100	120	170	200	230		
	1.25	Sample_129,99	Screen	Opening	26.670	22.400	19.200	16.000	13.550	6 520	7.930	6.350	5.613	3 963	3.360	2.794	2.002	1.700	1.400	1.000	710	. 600	.425	.355	. 250	.212	.150	.125	106	:075	.063	Totala	Gain or loss
	Project 1008 Project 100NS Location Sample Remarks CORE 58	Reight of		•	-6.67	-5.60	_	-4.00	-3.68	27.50	-3.00	-2.65	-2.50	2,50	-1.75	-1.50	-1.00	-0.75	-0.50	0.0	2 5	40.7	+1.25	+1.50	1 1.72	+2.25	+2.59	+3.00	+3.25	+3:73	2 1	- 4-	_
	Date May 1979	يو	Cumular (vol	Per Cent Passing	Ī																												
	Ā	, Date	Steves	Cumulative Per Cent					5.21	8.8	25.65	36.76	43.02		57.49	62.73	77 67	98.48	73.28	76.98	80.13		82,7/	84.25	85.21	9414	86.07	88.43		93.26	95,88	100.00	
A:SILYS18	,	S OF SAND	Recalication SI						5.21	777	57 0	11:11	6,26		73.5	10.24	2.0	21/8	4.84	3.69	3.15		2.64	1.68	48.0	82.7	9.86	2.35		4.83	2,62	4.12	
ᇤ		, ,,	- 12	1	1	1						ŀ		- 1					\vdash		4	1	- 1	1 1		+-		_			4	H	+
FRC SPIDERE	Collected by	IEVE ANALYSIS Analyzed by	2	Grams	1		-		5.95	4.33	10.80	12.69	7.15	+	- 78-0T	11.70	53,	25.0	5.53	4.22	3.60		3.02	1.69	01	74:4	96.0	2.69		5.52	2.99	4.71	116.23
CERC SPIDLIENT ANALYSIS	. 11	E .		1	73/6	2/8	3/4	5/8		+	5/16 10 80	+	3 1/2 7.15	1	+		10	+			+		70.5		50	+	H	+		200 5.52	+	P.10 4,71	
CENC SPIDLEST	Projections LAKE FRIE - ONIO Locations LAKE FRIE - ONIO Location Cone & 757 750 cm			Grams	7776 1 057 76	1	+		1/2	+	3/3	+	3 1/2	1	+	,	+	12	14	18	20	30	+	4.5	+	20	90	+	1-0	-	230	P.in	Total9 114.23

CERC SPUDIENT ANALYSIS

CERC SPIDILIENT ANALYSIS

CERC CN 0011
Project TCONS LAKE ERIE - 0HIO
Location/Sample No.
Recarra CORE 71 0 - 7 cm

SIEVE AMALY

Weight of Sample 188.62

Wesh Number

Screen Opening

s Date

Cumula	Par C																																						
670.5	Cumulactive Per Cent									191	2.63	4.52		9.88		16.71		26.31		42.10		60.35		74.42		81.82		84.85		88 49		92.38		97.47		98.83		99.16	100.00
Retained on Sleves	Per Cent									1,61	1.02	1.89		5,36		6.82		9,60		15,79		18.25		14.08		7.39		3.04		3.64		3,89		5.09		1.36		0.33	0.84
	Grams									1.90	1.20	2.23		6, 32.		8.05		11.32		18.62		21.52		16,60		8.72		3.58		4.29		4.59		9.00		1.60		0.39	0.99
U.S.	Kesh	W local	1 3/64	3/8	3/4	۶/۶	1/2	1/16	3/8	5/16	1/4	3 1/2	ব	5	. 9	7	. 8	10	12	14	16	18	20	25	30	35	7.0	45	50	90	70	80	100	120	1-0	170	200	230	Pan
Screen	Opening		26.670	22.400	19.200	16.000	13.550	11.100	9.520	7.930	6.350	5.613	6.760	3.962	3, 360	2.794	2.362	2.000	1.700	1.400	1.180	1.000	.850	.710	009.	. 500	.425	.355	300	. 250	.212	.130	.150	.135	.106	060.	.075	.063	0.000
	•		-6.67	-5.60	-4.80	50.4-	-3,68	-3.50	-3.25	-3.00	-2.65	-2.50	-2.25	-2.00	-1.75	-1.50	-1.25	-1.00	-0.75	-0.50	-0.25	.00.0	+0.25	1+0.50	+0.75	+1.00	+1.25	+1.50	+1.75	+2.00	+2.25	+2.50	+2.75	+3.00	+3.25	+3.50	+3.75	છ. 7	
Cumulative	Per Cent	An i de la																																					
5060	Cunulactive Per Cent				4.42	17.29	20.21	24.93	30.10	34.70	. 37.22	39.19		42.08		43.94		45.89		47.91		51.84		61.95		74.37		83.04		90.67		96.53		98.46		98.99		99,25	100.00
Retained on Sleven	Per Cont				4.42	12.87	2,92	4.73	5.17	4.60	2.52	1.97		2.89		1.86		1.94		2.02		3,93		10.11		12.41		8.68		7.62		5,86		1.93		0.53		0.27	0.75
_	Grams				8.30	24.10	5.48	8,88	9.71	8.64	4.73	3,70		5.43		3.50		3.65		3.80		7.39		18.99		23,32		16,30		14.32		11.01		3,62		1.00		05.0	1.40

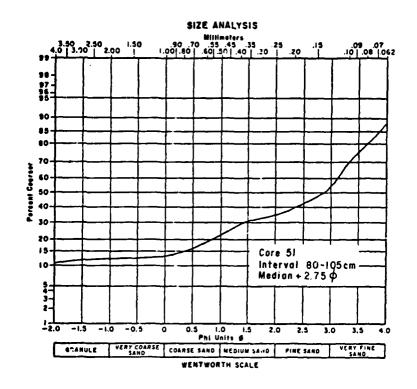
	Date Ney 1978		" Date	lov) reliming	Cunulative Per Cent Per Cent Passing			9.17	13.79	20.01		27.22	41	34,83		, wz	15		56	7.1		7	88			63	99		70	57			58	
I ASSLYSIS	ьу		IS OF SAND by	Retained on Steves	Per Cent Gunu		-	9.17		-6.23 20	-	5.10	ŀ		+	9.84	6.48 51.15		6.40 57.56	8,16 65.71	+	72.4	4,60 77.89	1 77 1 79 66		0.97 80.63	3.03 83.66		8.42 92.07	6,50 98,57	76 0	4.04	0,17 92,58	
CINC SPULIENT ANALYSIS	Collected		SIEVE ANALYSIS OF Analyzed by	RE	Grams			10.89	5,49	2.40		2.50	11.7	1.68		77.79	7.70		7.61	69'6	180	377	5.47	2 10		- इसर	3.60		80.01	7.72	-	7,1	0.20	0.50
	LAKE ERIE - OHIO	30 - 35 cm	1,25 Br.	3 II	Mesh	79/1	7/8	3/4	5/8	1/2	7/16	5/16	1/1	3 1/2	. 7	^		8	2 2	14	16	20	25	3,30	02	45	9	70	103	120	0,71	200	230	Pan
	1CONS LAK	Staple No.	Sample 119.25	Serve	Opening KM	26. 670	22.400	19.200	16.000	13.550	11.100	7,930	6.350	5.613	4.760	3.362	2.794	2.362	2.000	1.400	1.180	.850	. 210	909	.425	.355	. 250	.212	150	.135	901.	5,0.	.063	000.0
	CERC CA	Location/Sapple No Regards 50	Veight of		•	19 9-	-3.60	-4.80	-4.00	-3.68	- 2	-3.00	-2.65	-2.50	-2.25	-2.00	-1.50	-1.25	1.8	-0.50	-0.25	+0.25	+0.50	\$ \$ \$	+1.25	+1.50	+2.00	+2.25	+2.59	+3.00	+3.25	+3.75	2 7	
	Date Hay 1979		Date	Cumulativel	Per Cent Passing																													
٠			•	Steves	Cunulacive Per Cent					3.64		98.4	7.36	8.68			22.75		31.49	45.74	61.12		76.82	83.51		85.51	87.73	27.00	200	91.63	95.36		97.63	100.00
CENC REDIVENT ANALYSIS	by		SIS OF SAND	Retained on Si	Per Cant					3.64		***************************************	2,89	132	1		7,64		8.14	14.25	17.38		13.70	69'9		7.03	2.16	02, 6		1.20	1.73		2.27	
erse brive	Collected		SIEVE AMALYSIS Analyzed by_		Grams					06.90			5.48	2.50			14.49		16.29	27.05	32.99		26.00	12.70	3,	777	4.10	513		60.0	3.29	5.	200	, .
	- 14	54 cm	80	U.S.	Mesh	1 3/64	7/8	3/4	5/3	1/2	3/8	5/16	1/2	3 1/2	.,	9	,	80 9			18	20	57	3	0,7	50	09	80,0	100	07.1	170	200	630	
	E E	(=:]										90	ا ورا	76	29	3 8	ΙÌ	 शह										$ \ $			
·	1 1	Recaris CORE 87 50 - 54 of	Weight of Sample 191,80	Screen	Opening	26.670	22.400	19.200	16.000	13.550	2005	7.930	6.350	5.613	1 967	3.360	2.194	2.7	1.700	1.100	1.000	9.50	01/10	500	575	300	35.	130	150	.106	060	2,0	000	

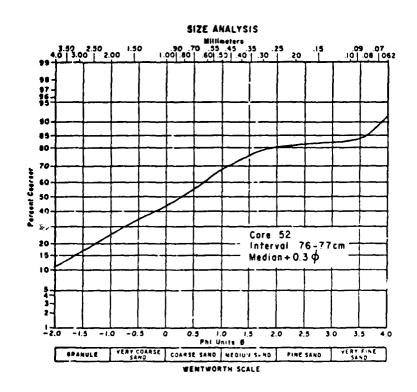
Froject T	2	LAKE ERIE - OHIO	Collected by			_							
		İ	1	d by	4	Date May 1979	Project	, , ,	OTRO - ZIEZ ZXVI	Collected by	by		Date May 1979
Location/Sanate	50	- 60 cm					Recarris CORE 90	CORE 90 75	5 - 80				
ğ	Seight of Sauple 160.69	•	SIEVE ANALYSIS OF Analyzed by	YSIS OF SAND	Ã.	r Date	Weight of	Weight of Sample, 142,10	i i	SIEVE AUALYSIS OF Analyzed by	SIS OF SAND	Ā	2 Date
ŀ	Coroso	-		Retained on Steves	16705	Cumulartue		Screen	1 0.5.		Retained on Sleve	STEVES.	Cumulative
	Opening Pri	Mesh	Grams	Per Cont	Cumulative Per Cent	Per Cent Passing	•	Opening PN	Mesh	Grams	Per Cont	Cumulative Per Cent	Per Cent Passing
-6 67	26,670	1 3/64					-6.67	26.670	1 3/64				
-5.60	2200	7/8					-5.60	22.400	7/8				-
-4.80	19.200	3/4					-4.80	19.200	3/4				
50.7-	16.500	£/5					00.7-	10.000	3/8				
1.56	11.100	3//					-3.50	11.100	7/16				
1	9.520	3/3	1.45	0.91	0.91		- 3.25	9.520	3/8		1. 0	.,	
	7.930	5/15					-3.00	7.930	5,16	7.00	1,51	2,74	
- -	6.350	1/4	10.10	2.57	3.48		-2.50	5.613	3 1/2	3.70	2.61	4.96	
1	7	,					-2.25	4.760	٠7				
- 2.00	3.962	5	14,00	3.77	14.18		-2.00	3.962	5	16.50	11.67	16.63	
-1.75	3.300	9		1	00 00		-1.73	2.794	,	27.05	19.13	35.75	
1	1.75	00	23.31	14:/5	78.30		-1.25	2,362	8				
-1.00	2.000	:0	25.65	16.06	44,97		-1.00	2.000	01	26.02	18.40	54.15	
	1.760	12					-0.75	1.700	12	31 18	22.05	76.20	-
-0.50	1.400	14	31.45	19.70	99.69		-0.25	1.180	16	3111	22744		
1	1.000	80	23.90	14,97	79.62		00.0	1.000	18	20.90	14.78	90,98	
	.850	20			30 30		+0.25	710	25	5.17	4.04	95.02	
+0.7.0+	009	000	10.11	6,22	22.23		+0.75	009.	30				
©; ∓	. 500	35	1,69	1.05	87.01		+1.60	. 500	35	0.50	0.35	35.3/	
+	536	2.7	3	72.0	37 78		100.1	.355	45	0.13	0.09	95.47	
- 1.2	300	20	7				+1.75	. 300	20			,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,	-
+2.00	. 250	90	2.08	1.30	88.66		+2.00	.250	99	0.41	0.29	33.78	
	. 212	0/0	0.8	2 3	93 98		42.50	.180	80	2.43	1.72	97.48	
1	.150	100	2	,,,,,	22.28		+2.75	.150	100				
+3.00	.135	120	1.00	0,63	94.61		+3.00	.125	120	2.50	1.77	99.24	
35.14	000	170	133	85 7	8		+3.50	060	170	0.39	0.78	99.52	
1, 75	5/0:	200	,,,,,	SP 4 E	2772		+1.75	.075	200				
ου: 7	. 063	230	0.37	0.23	99.42		00.7	.063	230	0.19	0.13	99.65	1
#	0.000	P.in	00.1	0.63	100.05			00.00	Lea -	×*.0	0.00	100.00	
			150							17 17			L

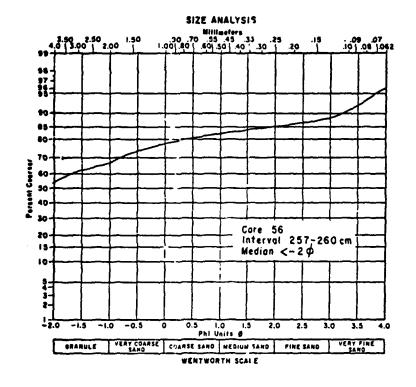
CERC SEDUIENT ANALYSIS

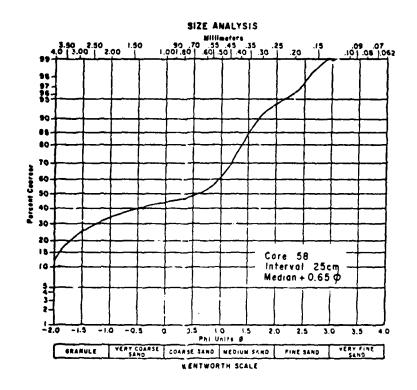
CERC CH	0094		Collected by		Date May 1979
Project	ICONS LA	KE ERIE - OL	110		
	Saaple No	ERIE, PA			
Rewarks_	CORE 94 1	00 - 110 cm			·
		•	SIEVE ANALYSIS OF SAND	•	· i
Weight o	f Sample_155		Analyzed by		Date
	Screen	U.S.	Retained on		Cumulative
	l Openine l	Mesh	1 1 Cash	Cumulative	Per Cent

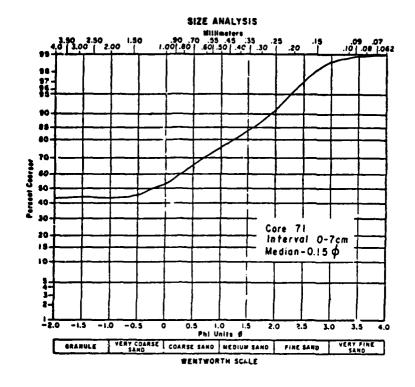
	Screen	U.S.		Retained on S		Cumulative
•	Opening 121	Mesh Number	Grams	Per Cont	Cumulative Per Cent	Pen Cent Passing
-6.67	26.670	1 3/64	 	 	 	
-5.60	22.400	7/8		 	 	
-4.80	19,200	3/4		 	 	
-4.00	16,000	5/8	 	ļ		
-3.68	13.550	1/2	 	 	}	
-3.50	11.100	7/16	3.16	2,33	2,33	
-3.25	9.520	3/8	3,23	2.09	4.42	·
-3.00	7,930	5/16	6.60	4.27	4.69	
-2.63	6.350	1/4	10.89	7.04_	15.74	
-2.50	5.513	3 1/2	4,81	3,11	18.84	
-2.25	4,760	3 27 2	4,01			
-2.00	3,962	5	12.70	8,21	27.06	
-1.75	3,360	6	12.70	- 9.41	 	
-1.50	2.794	7	12,59	8.14	35.20	
-1.25	2.362	8	1-14-17		 	
-1.00	2.000	10	12.70	8.21	43.42	
-0.75	1.700	12	 	B. 21	<u>#3.#/</u>	· · · · · ·
-0.50	1,400	14	20.55	13.29	56.71	
-0.25	1.180	16		13.49	 	
0.00	1.000	18	17.88		68.27	
+0.25	.850	20	1		1 - 00-44	
+0.50	.710	2.5	8,99	5.81	74.08	
+0.75	.600	30			i	
+1.00	.300	35	2,93	1.89	75.98	
+1.25	.425	40	 		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	
+1.50	. 355	45	2.86	1.85	77.83	
+1.75	.300	50			1	
+2.00	.250	60	9.15	5.92	83.75	
+2.25	.212	70	7.0			
+2.50	.130	0.6	16.39	10.60	94.35	1
+2.75	.150	100				
+3.00	.125	120	6.95	4.49	98.84	1
+3.25	.106	140	1		1	1
+3.50	.090	170	0.85	0.55	99.39	 -
+3.75	075	200	+ ···			1
+4.00	.063	230	0.25	0.16	99.55	
	0.000	Pan	0.69	0.45	100.00	
	Totala		154.62		AVVAVV	
	in or lo		-0.53		<u> </u>	

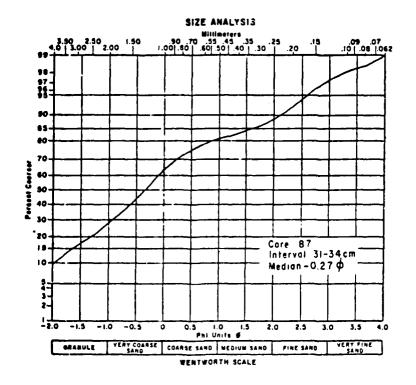


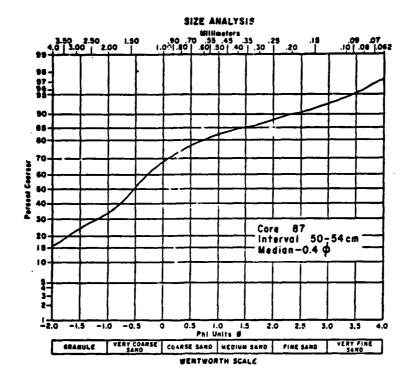


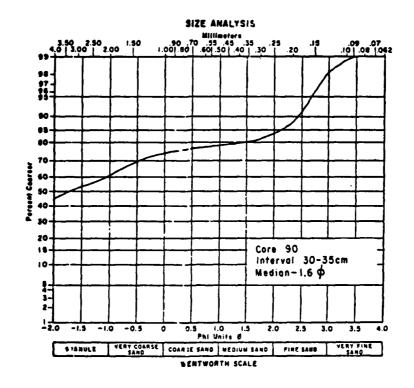




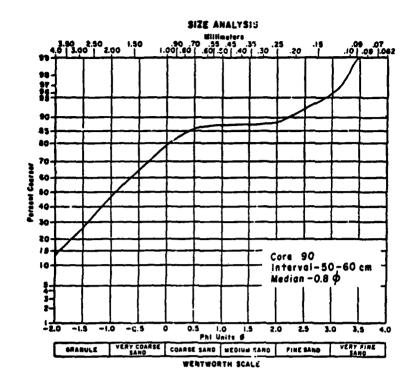


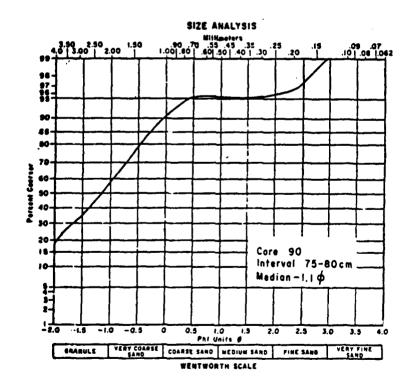


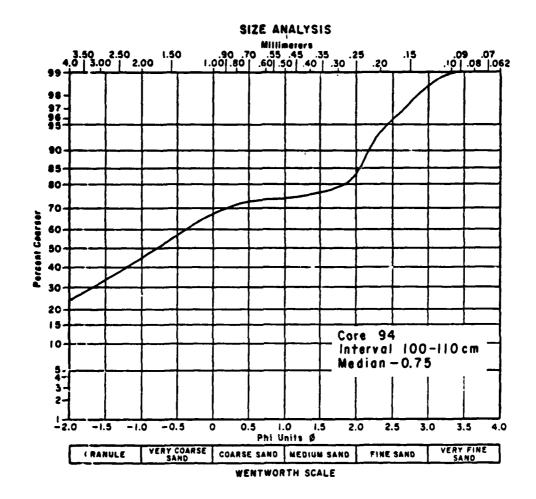




U







PIPET ANALYSIS

Core	(ca)	SAMO (pct)	SILT (pct)	CLAY (pct)	Core	IFTERVAL (CB)	SAMD (pct)	SILT (pct)	CLAY (pct)
- 44	69 to 70	18.38	74.92	6.7					9.46
**	110 te 120	14.75	57.69	27.56	87	100 to 110	13.18	77.36	
	!				1	250 to 260 550 to 560	5.51 4.82	80.67 79.90	13.82 15.28
49	30 to 40	39.26	40.71	20.03		f .			
	200 to 210 330 to 340	1.75 0.31	67.8 87.55	30.45	54	60 to 70	10.66	60.12	29.22
	400 to 410	2.17	90.12	11.94 7.64	1	200 to 210 400 to 410	3.41 3.85	73.26 43.94	23.33 52.21
	1				ł :				
55	30 to 40 190 to 195	1.79	63.1	35.11	89	330 to 340	48.41	36.00	15.59
	380 to 390	59.80 11.57	29.49 67.66	10.71 20.77	!	390 to 392	0.46	66.71	32.83
	410 to 420	17.53	42.37	40.10	90	335 to 337	51.64	34.88	13.48
57	170 to 180	70.57		6.45		500 to 510	2.43	80.95	16.62
21	270 to 280	50.37	22.98 42.2	7.43	91	105 to 107	52.30	30.32	17.38
	370 to 380	37.04	53.51	9.45	l	300 to 310	4.36	77.19	18.45
58	1				95	15 to 25	1.52	54.62	43.86
	440 to 450	83.06	16.1	0.84		; 80 to 90	0.92	71.35	27.73
59	247 to 248	24.66	61.89	13.45	1	280 to 290	2.07	69.95	27.98
60	360 to 370	62.51	30.37	7.12		390 to 392	30.29	57.72	11.99
	470 to 480	35.88	52.99	11.13	i	435 to 445	11.60	40.26	48.14
	560 to 570	18.92	42.12	38.96	96	0 to 5	13.49	62.67	23.84
61	40 to 50	71.0	22.47	6.53		66 to 67	0.41	77.56	22.03
	100 to 102	69.7	23.5	6.7	ŀ	130 to 140 200 to 210	0.86 0.98	63.39 57.34	35.75 41.68
	190 to 200	37.6	50.9	11.5	ł	250 to 260	13.59	46.63	39.78
	280 to 282	9.75	74.82	15.53	l .	285 to 295	10.92	51.93	37.51
	310 to 120	27.4	36.5	36.1	j	390 to 400	15.77	38.14	46.09
62	70 to #0	5.2	57.3	37.5		5.20	10.52	38.35	51.13
	160 to 170	26.8	58.4	14.8	<u> </u>	480 to 490	4.08	26.14	69.78
	260 to 270 320 to 330	1.6 61.62	53.5 25.43	44.9 12.95	98	60 to 70	2.65	75.74	21.61
					!	155 to 165	5.76	78.97	15.27
63	50 to 60 220 to 222	8.38 69.7	60.91 17.9	30.71 12.4	1	270-to 280	0.66	68.64	30.70
	340 to 350	2.99	67.86	29.15	· .	315 to 325	19.20	44.45	36.35
	450 to 460	1.30	48.94	49.76	99	30 to 40	2.12	75.88	22.00
66	60 to 70	2.21	61.05	36.74	ĺ	100 to 110 250 to 260	0.26	68.43	31.31
-	150 to 160	29.01	46.55	24.44		310 to 320	16.85 39.80 -	42.88 57.72	40.27 2.48
	224 to 226	8.02	\$7.65	34.33					
	350 to 360	5.31	\$9.15	35.54	101	553 to 556	72	25	2
	450 to 460	3.83	52.89	43.28	102	O to 🕹	25	51	24
67	40 to 30	61.59	25.01	13.4	ŀ	138 to 141	52	47	1
	170 to 180	51.7	30.7	17.6		252 to 255 332 to 335	36 2	62 74	2 24
	260 to 270	1.56	49.09	49.35	ł	340 to 343	ŝ	36	61
	360 to 370 460 to 470	1.1 1.07	54.4 38.57	44.5			-		2
				60.36	103	140 to 143 161 to 163	6 22	92 75	i
69	250 to 260 340 to 350	72.86	17.40	9.74	l	181 to 184	2	94	Ā
	440 to 450	53.29 2.66	34.98 52.68	11.73 44.56	1	207 to 209	3	94	3
	540 to 550	1.37	49.2	49.43	•	21.7 to 225	1	81	8
72	10 to 20	7.96	80.21	11.83	104	5 to 10	Q	60	40
	65 to 66	18.43	61.62	19.95	ł	64 to 69	22	62	16
	150 to 160	11.77	62.02	26.21		95 to 105	21	64	15 21
	"40 to 350	0.34	51.82	47.84	1	136 to 141 149 to 150	5 40	74 55	<u>د</u> ند 5
74	20 to 30	5.7	29.0	65.3	l .	202 to 206	<u> </u>	80	16
	285 to 295	5.7	26.0	68.3	1	300 to 305	T	63	37
	470 to 480	1.7	47.6	50.7	105	20 to 25	54	36	10
77	100 to 110	14.42	54.96	30.62		30 to 80	58	39	3
	525 to 535	18.01	41.08	40.91	!	99 to 104	15	81	4
79	60 to 63	32.05	49.8	18.15	i	160 to 165	7	76	17
	105 to 115	21.66	40.1	38.24	l	262 to 265 275 to 280	7	76 66	17 34
	300 to 310	23.26	39.66	37.08	ĺ	327 to 330	Ö	76	24
81	200 to 210	1.1	73.5	25.4	٠		-	4.9	2
	310 to 320	0.4	45.8	53.8	106	12 to 15 81 to 94	55 20	43 73	;
	400 to 410	11.8	37.7	50.5	Į.	171 to 173	6	90	Á
83	40 to 45	20.95	52.54	26.51	j	354 to 359	ž	52	41
	100 to 110	2.53	71.18	26.29		426 to 429	9	49	42
	475 to 485	2.05	60.50	37.45	107	0 to 5	20	65	15
84	250 to 260	30.35	67.84	1.81	i -~"	13 to 36	42 .	57	1
	335 to 345	16.60	40.16	43.24	1	70 to 80	65	35	0
85	100 to 110	11.79	71.91	16.3		131 to 135	76 57	24 42	0 1
	370 to 380	6.18	77.64	16.18	l	135 to 139 362 to 365	71	82	17
	405 to 415	15.30	37.85	46.85]	369 to 372	4	49	47
86	90 to 100	16.61	73.32	10.07	1				
	240 to 250	15.84	41.02	43.14					

¹Trace amount.

VI	SUA	L	AC(CUMUL	MOITA	TUBE	(VAT)
Core		(CE)		>4 ph1 (pct)	<4 phi (pct)	MEAN phi	STANDARD DEVIATION
101	130	to	60 140 200	84 82 91	16 19 9	3.9 3.8 3.6	0.7 0.6 0.6
103	0	to	5	98	2	2.8	0.7
106	7	to	10	95	5	3.1	0.7

APPENDIX C X-RAY DIFFRACTION ANALYSES

2		Mixed layer	×	_		*2			*						*		*	*	*		*	*
fraction (smear)		Kaolinite	×	×	×	×	×	×	×	×	*	×	*	*	×	×	×	*	×	×	. ×	*
um fract		Illite	×	×	×	×	×	×	×	×	×	×	×	×	×	×	×	×	×	×	×	×
Clay minerals <2		Chlorite- Vermiculite	×	×	×	×	×	×	×	*	×	×	×	×	×	×	×	×	×	×	×	×
Clay		Montmo- rillonite	Х		×					×			×				×				×	
	au .	Dolomite	43	18	33	43	21	23	84	99	32	30	87.	47	28	56	67	22	26	67	-3	55
	base line	Calcite	0	14	13	12		16	19	0	21	42	12	39	12	21	14	59	24	36	35	42
пш	tes above	Plagio- clase	32	20	07	28	16	24	31	39	41	24	04	41	47	55	. 67	27	56	34	54	38
Smear mounts <44 µm	k intensities	Potash feldspar	16	18	20	23	18	18	14	37	19	20	77	3.5	30	21	21	25	13	20	24	24
Smear m	Peak	Quartz (20.66 [°] 29)	8/	59	93	57	41	57	99	83	63	09	83	76	72	42	98	77	59	53	16	61
	Ratios	Kaolinite + Chlorite/ Quartz	0.45	1.54	0.40	0.74	2.27	1.63	0.58	0.24	1.00	0.93	0.24	0.29	0.51	1.93	0.21	1.80	0.78	0.57	0.12	0.82
	Rê	Mica/ Quartz	0.36	1.41	0.28	0.58	2.10	1.33	0.47	0.18	0.87	0.72	0.16	0.20	0.33	1.83	0.16	1.50	0.76	0.55	0.13	0.54
Material		glacial)	Muddy sand (P)	Gravelly clay (T)	Silt (P)	Silt (P)	Gravelly clay (T)	Clay (T)	Clay (T)	Silt (P)	Gravelly clay (T)	Gravelly clay (T)	Silt (P)	Silt (P)	Silt (P)	Gravelly clay (T)	S11t (P)	Clay (T)	Gravelly clay (T)	Clay (T)	Silt (P)	Gravelly clay (T)
Interval	(cm from	core)	105	410	99	250	315	75	260	78	510	320	07	175	370	077	110	250	320	07	250	450
Core	0		55	55	61	61	61	7.5	72	11	7.7	78	95	95	95	95	66	66	66	104	106	106

1X = present.
2* = minor amount present.
30ff scale.

APPENDIX D
ATTERBERG LIMITS

Core No.	Interval (cm from top of core)	Material (T = till; P = postglacial)	Liquid limit (pct)	Plasticity index
55	398 to 420	Gravelly clay (T)	28	10
60	560 to 580	Gravelly clay (T)	28	9
61	190 to 200	Silt (P)	0	0
61	280 to 290	Silt (P)	25	4
61	310 to 325	Gravelly clay (T)	25	7
72	70 to 77	Clay (T)	26	6
72	340 to 347	Clay (T)	31	11
74	10 to 25	Clay (T)	39	17
74	470 to 490	Clay (T)	33	13
77	200 to 220	Gravelly clay (T)	29	10
79	100 to 125	Gravelly clay (T)	27	11
7 9	300 to 320	Gravelly clay (T)	25	8
81	310 to 320	Clay (T)	31	11
81	400 to 410	Clay (T)	28	8
84	335 to 350	Gravelly clay (T)	0	0
85	200 to 210	Silt (P)	0	0
85	370 to 380	Silt (P)	29	11
85	400 to 420	Gravelly clay (T)	29	13
86	230 to 250	Gravelly clay (T)	29	11
95	430 to 445	Gravelly clay (T)	31	11
96	285 to 295	Gravelly clay (T)	22	6
96	480 to 490	Gravelly clay (T)	38	18
98	316 to 325	Gravelly clay (T)	27	10
99	240 to 255	Clay (T)	27	8
99	305 to 320	Gravelly clay (T)	22	6
105	130 to 140	Silt (P)	22	3
105	310 to 320	Silt (P)	33	26
105	370 to 380	Clay (T)	34	12
105	440 to 455	Clay (T)	28	9

APPENDIX E

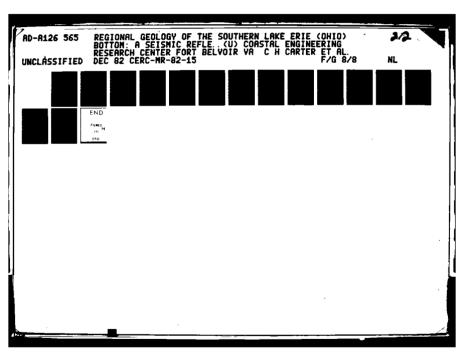
MOLLUSK IDENTIFICATIONS FROM CORES

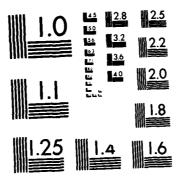
	Abundance ¹
Phylum - Mollusca	
Class - Bivalvia	
Order - Prionodesmacea	
Family - Unionidae	
Genus - Elliptio	
E. dilatatus sterkii	-
Subfamily - Lampsilinae	
Genus - <i>Lampsilis</i>	
L. radiata	-
Lampsilis radiata siliquoidea	-
Genus - Obliquaria	
Obliquaria reflexa	-
Order - Teleodesmacea	
Family - Sphaeriidae	
Genus - Sphaerium	
Sphaerium sp.	+
Class - Gastropoda	
Order - Ctenobranchiata	
Family - Amnicolidae	
Genus - Amnicola	
Amnicola integra Amnicola lacustris	+
	+
Amnicola leightoni Amnicola limosa	+
Amnicola sp.	+
	т
Subfamily - Lithoglyphinae Genus - Somatogyrus	
Somatogyrus subglobosus	_
Subfamily - Bulimnae	_
Genus - Bulimus	
Bulimus tentaculatus	0
operculum from Bulimus tentaculatus	_
Family - Pleuroceridae	
Genus - Pleurocera	
Pleurocera acutum	0
Pleurocera sp.	0
Genus - Goniobasis	· ·
Goniobasis liviscens	0
Goniobasis haldemani	0
Family - Valvatidae	•
Genus - Valvata	
Valvata sincera	+
Valvata tricarinata	+
ramily - Pomatiopsidae	
Genus - Pomatiopsis	
Pomatiopsis lapidaria	-

¹Abundance: + is abundant; 0 is common; - is rare.

	Abundancel
Order - Pulmonata	
Family - Lymnaeidae	
Genus - Fossaria	
Fossaria parva	_
Fossaria modicella rustica	_
Fossaria cf. exigua	-
Family - Planorbidae	
Genus - Gyraulus	
Gyraulus paryus	-
Subfamily - Planorbulinae	
Genus - Promenetus	
Promenetus exacuous	-
Promenetus umbilicatellus	_
Family - Physidae	
Genus - Physa	
Physa gyrina	_
Family - Valloniidae	
Genus - Vallonia	
Vallonia costata	-

 $^{^{1}}$ Abundance: + is abundant; 0 is common; - is rare.





MICROCOPY RESOLUTION TEST CHART NATIONAL BUREAU OF STANDARDS-1963-A

APPENDIX F

SEDIMENT THICKNESS DATA FROM SEISMIC RECORDS

Seismic velocities:

water v = 1.54 meters per millisecond

postglacial sediments v = 1.3 meters per millisecond

 $\overline{v} = 1.8$ meters per millisecond

Water depth is survey water depth (about 174.3 meters), which is 1 meter above low water datum (IGLD, 1955).

- a. Postglacial sediment and till thicknesses were measured at two confidence levels. Those followed by an "a" (the first confidence level) are those in which the reflector was well defined on the seismic record; unmarked thicknesses are those in which the reflector was not well defined on the seismic record yet whose position could be reasonably well inferred from the overall geologic setting and character of the reflector.
- b. The letter "b" indicates thickness data not available because (1) both postglacial sediment and till thicknesses could not be measured because of difficulty in interpreting the seismic records (in this case water depth was not measured), or (2) the seismic reflection record for the fix is missing.

All thicknesses in meters

- T = trace of material, <0.5 meter, but enough to change surface echo character.
- x = thickness not measurable because of lack of lower reflector.

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
369	8.5	0a	0a	461	to 475	1	b
370	8.5	0a	0a				
371	11.2	0a	0a	476	14.8	0	0
372	13.6	2.0a	0a	477	15.0	0	0
373	14.1	2.5a	3.5a	478	to 482	1	Ь
374	15.4	4.5a	5.0a				
		,,,,,,		483	13.6	0	0
375	to 386		b	484	14.5	2.5	x
387	10.0	0	0	485	14.9	4.0	x
388	10.1	Ō	Ō	486	15.0	4.0	x
		· ·		487	15.1	5.0	x
389	to 422		Ъ	488	15.5	6.0	x
423	15.7	4.5a	5.0a	489	16.4	7.5a	x
424	15.6	2.0a	6.0a	490	16.9	8.5a	x
425	15.0	2.0a	4.0a	491	16.6	9.0a	x
426	14.5	2.5	0	492	17.3	10.5a	x
427	14.1	1.0	0	493	17.7	10.5a	x
428	13.3	0a	0 0a	494	17.8	11.5a	x
429		0a 0a		495	17.4	12.5a	x
	11.5		0a	496	18.6	12.5a	x
430	8.9	0a	0a	497	19.0	12.5a	x
431	12.4	0a	0a	498	20.5	14.0a	x
432	13.9	0a	0a	499	20.9	15.0a	x
433	14.7	0	4.5	500	20.5	16.0a	×
434	15.0	1.0a	6.5a	501	20.6	16.0	x
435	15.3	3.0a	5.0a	502	21.8	17.0	x
436	16.8	4.5a	4.5a	503	21.8	17.0	x
437	16.0	5.0a	9.5a	504	21.8	17.0	x
438	16.6	5.5a	13.0a	505	22.4	17.0	×
439	16.8	6.0a	13.5a	506	22.6	17.0	x
440	16.6	7.5a	12.5a	507	22.7	17.5	x
441	17.0	7.5a	12.0a	508	22.6	17.0	x
442	17.6	7.0a	12.0a	509	22.1	17.0	x
443	16.5	5.5a	10.0a	510	21.8	17.0	
444	16.9	4.0a	10.0a	710	21.0	17.0	x
445	16.1	3.0a	9.0a	511	to 531	į	b
446	15.3	3.0a	6.0a	532	15.6	6.5	2.0
447	14.8	2.5a	3.0a	533	16.5		2.0
448	14.2	1.0a	1.5a			8.0a	2.0a
449	12.2	0a	0a	534	17.0	8.0a	1.5a
450	10.4	0a	0a	535	16.4	6.5a	3.0a
451	12.2	0a	0a	536	16.2	5.5a	2.0a
452	14.7	0	0	537	15.2	3.5a	2.0a
453	15.0	1.5	x	538	14.8	3.5a	1.0a
454	16.0	3.5	x	539	14.5	3.0a	1.0a
455	15.5	5.0	6.0	540	12.1	0a	0a
456	16.2	6.0	7.5	541	11.2	0a	0a
457	16.8	8.0	9.0	542	8.0	0a	0 a
458	16.9	8.5	х	543	8.8	0a	0 a
459	16.3	8.5	x	544	11.7	0a	0a
460	16.7	8.5	x	545	12.5	1.0a	0 a
,00	±0 • r	J.J	^	546	13.8	1.5a	0.5a

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
547	15.2	1.5a	2.0a	597	21.5	7.5a	21.5a
548	15.2	2.0a	1.5a	598	21.6	9.5a	21.5
549	15.5	2.0a	2.0a	599	21.6	12.0a	19.5
550	15.8	2.0a	4.5a	600	22.0	12.5	25.0
551	15.8	2.0a	4.5a	601	22.3	13.5a	23.5
552	16.8	2.0a	5.5a	602	22.1	13.5a	23.0
553	17.3	2.0a	9.0a	603	21.9	13.0a	21.0
554	17.8	2.0a	8.0a	604	21.6	12.0a	18.5
555	16.0	2.0a	5.0a	605	21.6	11.0a	16.0
556	15.3	2.0a	2.5a	606	21.4	10.5a	
557	15.0	1.5a	2.0a	607	21.4	9.0a	14.5a
558	14.7	0a	0a	608	20.8		14.0a
559	12.0	0a	0a	609	20.5	8.5a	10.5a
560	10.7	0a 0a	0a	610		8.5a	8.5a
561	8.9	0a	0a	1	19.6	8.0	8.5
562	10.5	0a 0a	0a	611	19.3	7.5	7.0
563	11.6	0a 0a		612	19.3	7.5	3.5
564	13.3	0.5	0a 0	613	19.1	7.5	0
565	13.9			614	18.5	6.0	0
		2.0a	1.0	615	17.4	5.5	0
566	14.6	2.0a	1.5	616	16.8	4.5	0
567	15.0	2.5a	2.0	617	15.5	3.0	0
568	16.2	2.5a	2.5a	618	14.1	2.5	0
569	16.5	2.5	2.5	619	13.9	1.5	0
570	17.3	2.5	5.5	620	13.4	0a	0a
571	17.7	2.5	7.5	621	11.2	T	0
572	18.0	2.5	9.5	622	10.0	T	0
573	18.9	4.0a	8.5a	623	10.7	T	0
574	18.6	3.0a	10.0a	624	13.0	T	0
575	17.6	2.5a	6.5a	625	to 636	1)
576	17.3	3.0a	5.0a	ł			
577	16.2	2.0a	3.5a	637	13.0	T	0
578	15.3	1.5a	3.5a	638	11.3	0a	0a
579	15.3	1.5a	1.5a	639	9.0	0a	0a
580	14.2	1.0a	0a	640	8.9	0a	0a
581	12.2	0a	0a	641	10.2	0a	0a
582	10.0	0a	0a	642	11.3	0a	0a
583	8.5	0 <i>a</i>	0a	643	11.5	0a	0 a
584	8.9	0a	0a	644	12.4	T	x
585	11.7	0a	0a	645	to 656	ł	
586	13.1	0a	0a	1	20 030	•	,
587	15.2	2 . 5a	0a	657	13.7	1.0a	0
588	15.6	3.0a	0a	658	11.6	2.0a	0
589	17.6	3.5a	3.0a	659	8.9	2.0a	0
590	18.8	4.0a	5.0a	660	8.8	2.0a	0
591	18.2	4.5a	7.5a	661	9.7	1.5a	0
592	19.1	5.5a	8.0a	662	10.6	1.5a	0
593	20.1	6.0a	10.0a	663	10.5	2.0a	0
594	20.7	6.5a	11.0a	664	10.8	1.5a	0
595	20.9	7.0a	13.0a	665	11.2	1.5a	0
596	21.2	8.0a	14.0a	666	11.7	1.5a	0

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
667	10.8	3.5a	0	745	17.9	4.0a	9.0
668	10.4	5.0a	x	746	17.5	3.5a	10.0
669	10.8	4.5a	x	747	17.4	2.5a	10.5
670	11.8	4.0a	x	748	17.1	T	13.0
671	12.9	3 .5 a	x	749	16.6	T	13.0
672	13.0	4.0a	x	750	16.6	T	12.5
673	12.3	5.0a	x	751	16.2	Ť	11.5
674	13.0	5.0a	x	752	14.9	T	11.5
675	16.3	4.0a	Y	753	15.8	0	10.5
			_	754	14.7	Ō	9.5
0/0	to 705	· ·	•	755	14.4	0	8.0
706	9.6	1.5a	x	756	14.4	0	7.0
707	10.5	1.5a	x	757	12.7	Ö	6.0
708	11.0	1.5a	x	758	11.8	0	5.5
709	11.0	1.5a	x	759	11.6	Ö	4.5
710	11.4	2.5a	x	760	8.8	Ō	5.5
711	12.0	2.5a	x	761	7.0	Ō	5.5
712	12.5	2.0a	x	762	8.2	Ō	5.0
713	12.7	1.5a	x	763	9.7	0	x
714	13.4	1.5a	x	764	10.4	0	x
715	14.0	1.5a	x	765	11.3	Ö	x
716	14.3	2.0a	x	766	11.4	Ö	x
717	14.9	2.0a	x	767	10.9	ő	x
718	15.0	2.0a	x	768	12.5	Ö	x
719	15.0	1.5a	x	769	12.2	Ö	x
720	15.6	2.0a	x	770	13.3	Ō	x
721	14.9	4.5a	x	771	13.1	0	x
722	16.4	5.0a	x	772	13.5	0	x
723	16.2	6.0a	x	773	14.5	0	x
724	18.0	6.5a	x	774	13.8	0	x
725	18.5	8.5a	x	775	14.1	0	x
726	19.3	11.0a	x				
727	20.4	12.0a	x	776		i	o
728	20.6	12.5a	x	777	14.9	0	10.0
729	21.5	14.0a	x	778	14.4	0	8.5
730	21.6	15.0a	x	779	13.2	0	8.5
731	21.6	16.5a	x	780	13.2	0	7.5
732	21.8	17.5a	x	781	12.1	0	7.0
733	22.1	17.5a	x	782	11.3	0	6.0
734	21.8	17 . 5a	x	783	10.7	0	5.5
735	21.7	17.0a	x	784	9.0	0	4.0
736	21.4	16.5a	x	785	6.5	0	3.5
737	21.1	15.0a	x	786	8.4	0	3.0
738	20.7	14.0a	x	787	11.4	0	6.0
739	20.3	11.5a	x	788	12.8	0	6.0
740	19.8	10.0a	x	789	12.6	0	6.5
741	19.5	8.5a	x	700	to 791	1)
742	19.2	7.5a	x	'90	LO / 7I		,
743	19.0	6.5a	x	792	12.2	0	7.5
744	18.7	5.0a	9.5	793	13.2	0	6.5

794 13.2 0 9.5 843 17.3 0 x 795 13.6 0 12.0 844 17.6 0 x 796 14.6 0 14.5 845 18.0 0 x 797 14.4 0 0 x 846 18.5 0 x 798 14.6 0 x 848 17.4 T x 800 14.9 0 x 848 17.4 T x 800 14.9 0 x 849 16.9 0 x 801 15.5 0 15.0 850 16.6 0 x 802 15.7 0 16.0 851 16.6 0 x 803 16.7 0 16.5 852 16.6 0 x 804 17.3 0 16.0 853 16.6 0 x	Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
795 13.6 0 12.0 844 17.6 0 x 796 14.6 0 14.5 845 18.0 0 x 797 14.4 0 x 846 18.5 0 x 799 15.8 0 x 848 17.4 T x 801 15.5 0 15.0 850 16.6 0 x 801 15.5 0 15.0 850 16.6 0 x 802 15.7 0 16.5 852 16.6 0 x 804 17.3 0 17.0 853 16.6 0 x 806 18.0 0.5a 18.0 854 16.0 0 x 806 18.0 0.5a 19.0 855 16.4 0 x 807 18.3 T x 856 15.0 0 x 80	794	13.2	0	9.5	843	17.3	0	х
796 14.6 0 14.5 845 18.0 0 x 798 14.6 0 x 846 18.5 0 x 799 15.8 0 x 848 17.4 T x 800 14.9 0 x 848 17.4 T x 801 15.5 0 15.0 850 16.9 0 x 803 16.7 0 16.0 851 16.9 0 x 804 17.3 0 17.0 853 16.6 0 x 805 17.8 0.5a 18.0 854 16.0 0 x 806 18.0 0.5a 19.0 855 16.4 0 x 807 18.3 T x 856 15.0 0 x 808 18.7 1.0a 16.0 857 14.0 0 x 81								
797 14,4 0 x 866 18.5 0 x 798 14.6 0 x 847 18.0 T x 800 14.9 0 x 849 16.9 0 x 801 15.5 0 15.0 850 16.6 0 x 802 15.7 0 16.0 851 16.9 0 x 803 16.7 0 16.5 852 16.6 0 x 804 17.3 0 17.0 853 16.6 0 x 806 18.0 0.5a 18.0 854 16.0 0 x 807 18.3 T x 856 15.0 0 x 808 18.7 1.0a 19.0 855 16.4 0 x 809 b 856 15.0 0 x 810 19.0 1.0a </td <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td>								
798 14.6 0 x 847 18.0 T x 799 15.8 0 x 848 17.4 T x 800 14.9 0 x 849 16.9 0 x 801 15.5 0 15.0 850 16.6 0 x 803 16.7 0 16.5 852 16.6 0 x 804 17.3 0 17.0 853 16.6 0 x 805 17.8 0.5a 18.0 854 16.0 0 x 806 18.0 0.5a 19.0 855 16.4 0 x 807 18.3 T x 856 15.0 0 x 808 18.7 1.0a 19.0 857 14.0 0 x 810 19.0 1.0a 16.5 860 13.5 0 x <td< td=""><td></td><td></td><td></td><td></td><td></td><td></td><td></td><td>x</td></td<>								x
Total Color				x				
800 14.9 0 x 849 16.9 0 x 801 15.5 0 15.0 850 16.6 0 x 803 15.7 0 16.5 851 16.6 0 x 804 17.3 0 17.0 853 16.6 0 x 805 17.8 0.5a 18.0 854 16.6 0 x 806 18.0 0.5a 19.0 855 16.4 0 x 807 18.3 T x 856 15.0 0 x 808 18.7 1.0a 19.0 857 14.0 0 x 809 b 859 13.6 0 x 810 19.0 1.0a 16.5 860 13.5 0 x 811 19.1 0.5a 17.5 861 12.8 0 x 812 18.6 0.5a 17.5 862 12.3 0 x 812 18.6 <td></td> <td></td> <td></td> <td>x</td> <td></td> <td></td> <td></td> <td></td>				x				
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840 16.6 0 x 890 15.1 0 x 841 17.5 0 x 891 15.0 0 x								
841 17.5 0 x 891 15.0 0 X								x
					1 892	15.0	0	x

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
893	15.0	0	x	942	14.9	1.0a	23.0a
894	15.0	0	x	943	13.9	0	22.0
895	15.1	0	x	944	12.6	0a	18.5a
896	14.3	0	x	945	11.6	0a	10.5a
897	13.6	T	x	946	8.4	0a	0a
898	13.5	T	x	947	5.7	0a	0a
899	12.0	T	x	948	6.8	0a	0a
900	10.4	0	x	949	8.7	T	9.5
901	7.0	0	x	950	9.0	T	11.0
902	7.2	0	x	951	10.6	T	10.0a
903	10.6	T	x	952	10.7	0	15.0
904	12.2	T	x	953	11.4	0	17.5
905	12.8	T	x	954	12.1	0	20.0
906	10.9	0	x	955	13.6	1.5a	18.5
907	10.0	0	x	956	14.6	1.0a	17.5
908		ь	i	957	15.0	1.5a	20.0
				958	14.5	1.5a	21.0
909	8.3	T	x	959	15.0	1.0a	23.0
910	10.0	T	x	550	15.2	1.0a	23.0
911	11.6	T	x	961	15.8	1.0a	22.5
912	13.6	0	x	962	15.2	1.0a	22.5
913	13.9	0	x	963	15.0	1.0a	23.0
914	14.5	0	x	964	14.5	1.5a	21.5
915	14.8	0	x	965	14.0	1.5a	21.5
916	15.2	0	x	966	13.5	1.0a	20.0
917	15.6	0	x	967	13.0	1.5a	17.5
918	15.8	0	x	968	12.0	1.5a	17.5
919	15.6	0	x	969	11.3	1.0a	15.0a
920	15.8	0	x	970	10.6	1.0a	1.5a
921	15.6	0	x	971	5.8	0a	0a
922	15.4	0	x	972	6.8	0a	0a
923	15.4	0	x	973	3.01	1.0a	6.0a
924	15.5	0	x	974	11.8	1.5a	14.0a
925	15.8	0	x	975	12.8	1.5a	18.0
926	15.5	0	x	976	13.2	2.0a	19.0
927	15.5	0	. K	977	13.7	1.5a	19.5
928	15.5	0	26.5	978	14.4	1.5a	20.0
929	14.1	0	25.0	979	14.9	1.0a	x
930	12.7	0a	26.0a	980	15.4	1.0a	x
931	11.7	0 a	14.0a	981	15.2	1.0a	x
932	4.7	0a	0a	982	14.8	1.5a	20.5
933	10.9	0a	0a	983	14.2	1.5a	20.0
934	12.7	0a	16.0a	984	13.8	2.0a	18.0
935	14.0	1.0a	23.5a	985	13.0	2.0a	18.0
936	15.4	1.0a	25.0a	986	12.4	2.0a	16.0a
937	15.9	0a	26.5a	987	11.4	2.0a	12.0a
938	16.3	0a	25.5a	988	10.9	1.0a	8.0a
939	16.4	0a	25.5a	989	8.3	0a	0a
940	16.2	0a	25.5a	990	8.1	0	3.5
941	15.5	0.5a	24.5a	991	8.4	0	5.0

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
992	8.0	0	9.5	1045	10.9	0a	0a
993	7.5	0	10.5	1046	8.2	0a	0a
994	6.4	0	7.0	1047	11.8	0	0
995	5.1	0a	0a	1048	12.9	1.5a	0
996	7.7	1.0a	0	1049	13.4	1.5a	4.0a
997	to 1001	ł	•	1050	13.9	1.0a	7.0a
			,	1051	14.0	1.0a	11.5
1002	14.9	1.5a	x	1052	14.1	1.0a	13.5
1003	15.5	1.5a	x	1053	14.8	0.5a	15.0
1004	15.9	1.5a	x	1054	14.8	0.5a	18.0
1005	15.8	1.5a	x	1055	14.2	0.5a	17.5
1006	15.2	1.5a	x	1056	13.8	1.0a	15.0
1007	15.2	T	x	1057	13.6	0.5a	12.5a
1008	15.2	T	x	1058	13.0	1.0a	7.0a
1009	14.6	T	x	1059	12.0	0a	5.0a
1010	13.7	T	x	1060	10.9	0a	3.5a
1011	12.1	T	x	1061	10.1	0	0
1012	9.0	T	0a	1062	7.0	0	0
1013	9.2	0a	0a	1063	8.9	0	T
1014	9.5	0a	0 a	1064	10.0	0	x
1015	13.3	0a	0a	1065	10.3	0	x
1016	15.7	1.5a	3.0a	1066	11.5	0	x
1017	16.0	1.5a	8.0a	1067	11.8	0	x
1018	16.2	1.5a	11.5a	1068	11.9	0	14.5
1019	17.0	1.5a	15.0a	1069	12.5	0	20.0
1020	17.0	1.5a	16.5a	1070	12.9	0	20.0
1021	17.0	1.5a	15.5a	1C71	12.8	0	21.5
1022	17.0	1.0a	15.0a	1072	12.9	0	22.5
1023	16.5	1.0a	13.5a	1073	14.4	0	20.5
1024	16.0	1.5a	12.0a	1074	15.9	0	21.0
1025	15.4	1.5a	7.5a	1075	16.6	0	22.0
1026	15.2	1.5a	1.5a	1076	16.8	1.0a	21.0a
1027	14.5	0a	0a	1077	16.8	2.5	x
1028	11.8	0a	0a	1078	17.0	5.0	x
1029	10.4	0a	0a	1079	16.8	6.5	x
1030	11.1	0a	0a	1080	18.5	3.5	x
1031	12.5	0a	0a	1081	16.5	2.5	x
1032	14.2	1.5a	0a	1082	16.2	2.5	x
1033	14.5	1.5a	2.0a	1083	16.0	3.2	x
1034	15.0	1.0a	8.0a	1084	15.8	3.0	×
1035	15.5	1.0a	11.0a	1085	15.5	5.2	x
1036	15.9	1.0a	13.0a	1086	15.0	5.0	x
1037	16.0	1.0a	13.5a	1087	15.0	3.0	x
1038	16.2	1.0a	16.5a	1088	14.8	5.5	x
1039	16.2	1.0a	15.0a	1089	13.8	6.5	x
1040	16.0	1.0a	14.5a	1090	13.8	6.0	x
1041	15.2	1.0a	13.0a	1091	14.5	4.8	x
1042	14.8	1.0a	9.0a	1092	15.5	4.2	x
1043	14.2	1.0a	4.0a	1093	15.7	4.8	x
1044	13.8	1.0a	1.5a	1094	16.1	4.5	x

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
1095	16.2	3.5	Х	1147	13.5	5.0a	9.0
1096	16.0	3.2	x	1148	13.5	5.0a	9.5
1097	14.5	3.8	x	1149	13.5	4.0a	6.5
1098	13.2	5.0	x	1150	13.1	3.0a	6.0
1099	14.2	3.5	x	1151	12.8	2.0a	4.5a
1100	12.9	3.5	x	1152	12.2	1.5a	3.5a
	to 1105		 b		and 1154		5.5 a b
1106	11.3	0a	6.5a	i			
1107	11.8	0a 0a	9.5a	1155	8.5	0a	0a
	12.2			1156	8.6	0x	0
1108		0a	9.5a	1157	8.5	0a	0a
1109	12.3	0	x	1158	8.5	0a	0a
1110	12.6	0	x	1159	9.0	0a	0a
1111	12.5	0	x	1160	11.0	0a	0a
1112	12.2	0a	13.5a	1161	11.8	1.0a	1.5a
1113	11.7	0a	10.0a	1162	12.1	2.0a	2.0a
1114	11.3	0a	7.5a	1163	12.2	3.0a	5.0a
1115	10.2	0a	6.0a	1164	12.2	3.5a	5.0a
1116	9.1	0a	5.5a	1165	12.5	4.5a	7.0a
1117	7.9	0a	4.5a	1166	12.8	4.5a	8.0
1118	7.6	0a	3.5a	1167	12.8	5.0a	9.5
1119	7.9	0a	3.5a	1168	12.8	5.0a	9.0
1120	9.4	0a	5.0a	1169	12.5	4.5a	9.0
1121	10.4	0a	6.0a	1170	12.2	4.0a	7.0a
1122	11.2	0a	8.5a	1171	12.2	3.5a	5.0a
1123	12.1	0a	9 .5 a	1172	11.8	2.5a	3.5a
1124	13.3	0a	10.0a	1173	11.2	1.5a	1.0a
1125	13.4	0.5a	10.5a	1174	9.2	0a	0a
1126	13.5	1.5a	x	1175	9.0	0a	0a
1127	14.0	2.0a	x	1176	7.8	0a	0a
1128	14.2	3.5a	x				
1129	14.5	4.0a	x	1177		·	Ь
1130	14.0	2.5a	9.5a	1178	11.0	1.0a	0a
1131	13.5	2.0a	9.0a	1179	11.0	1.5a	0 a
1132	13.2	1.5a	10.0a	1180	10.8	1.0a	0 a
1133	13.2	1.5a	8.5a	1181	10.0	0.5a	0a
1134	13.0	1.5a	6.0a	1182	9.0	0.92	0
1135	12.5	1.0a	5.5a	1183	8.2	0	0
1136	12.1	1.0a	1.5a	1184	7.8	0	0
				1185	6.5	0	3.0
1137	and 1138	1	b	1186	6.0	0	2.0
1139	9.0	0a	0a	1187	7.2		
1140	9.4	0a	0a	1188	8.5	0	0
1141	11.5	Ta	0a			1.0	4.0
1142	12.2	1.0a	0 a	1189	9.0	1.0	4.0
1142	12.2	1.0a 2.0a	0	1190	9.0	1.5	2.0
				1191	9.5	1.0	x
1144	12.8	3.0a	4.0a	1192	9.0	1.0	x
1145	13.0	4.0a	6.5a	1193	8.5	0.5	x
1146		I	b	1194	8.0	1.0	×
				1195	7.5	1.5	x

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
1196	7.0	1.0	х	1246	15.5	2.5a	x
1197	14.3	3.0	x	1247	15.5	3.0a	x
1198	13.7	3.0	x	1248	15.6	3.0a	x
1199	12.8	3.5	x	1249	15.8	3.5a	x
1200	12.3	2.5	x	1250	15.5	3.5a	x
1201	13.9	1.0	x	1251	15.5	2.5a	x
1202	14.4	1.0a	x	1252	15.5	2.5a	x
1203	14.9	1.0a	25.0	1253	15.5	1.0a	25.0
1204	14.8	1.0a	25.5	1254	14.8	T	28.0
1205	15.1	0	25.5	1255	14.4	0.5a	25.5
1206	15.7	0	25.5	1256	13.3	1.0a	28.0
1207	15.2	1.0a	24.5	1257	13.4	0.5a	28.0
1208	15.2	1.5a	23.0	1258	14.3	0	27.0
1209	15.3	2.5a	23.0	1259	14.5	0	27.0
1210	15.3	2.5a	22.5	1260	14.5	0	27.0
1211	15.4	3.5a	22.0	1261	14.7	0	27.0
1212	15.6	3.5a	22.0	1262	14.5	0	27.0
1213	15.5	3.5a	x	1263	14.7	0	28.5
1214	15.4	4.5a	x	1264	14.7	0	27.5
1215	16.2	3.5a	21.5	1265	13.7	0.5a	27.5
1216	16.2	4.0a	21.5	1266	14.9	1.0a	28.0
1217	16.0	4.0a	22.0	1267	15.0	1.5a	27.0
1218	16.2	3.5a	21.0	1268	15.3	1.0a	26.0
1219	16.2	3.5a	21.0	1269	15.5	2.0a	25.5
1220	16.0	2.5a	22.0	1270	15.7	1.5a	26.0
1221	15.9	3.0a	23.0	1271	15.9	2.0a	25.0
1222	15.8	2.5a	23.5	1272	15.8	2.0a	24.5
1223	15.2	2.5a	23.0	1273	16.2	2.0a	25.0
1224	15.9	1.5a	24.0	1274	16.3	2.5a	22.5
1225	15.6	1.5a	24.0	1275	15.3	2.5a	25.0
1226	15.6	1.5a	24.5	1276	15.5	2.0a	25.0
1227	15.6	1.0a	23.5	1277	15.6	2.0a	25.0
1228	15.0	0.5a	25.0	1278	15.5	1.5a	25.5
1229	14.9	0.5a	24.5	1279	15.6	1.5a	25.0
1230	15.0	0.5a	26.0	1280	14.8	1.5a	27.0
1231	15.1	0	26.0	1281	14.6	1.0a	26.5
1232	14.9	0	26.5	1282	14.1	1.0a	26.5
1233	14.7	0	26.5	1283	14.0	1.0a	28.0
1234	14.8	0	28.0	1284	13.9	0.5a	28.5
1235	14.2	0	26.0	1285	14.3	0 a	28.0
1236	14.1	0	25.0	1286	14.0	0a	28.0
1237	14.3	0	26.5	1287	14.7	0a	27.5
1238	13.6	0	26.5	1288	14.9	0a	27.5
1239	13.3	1.0a	27.5	1289	14.2	0a	27.5
1240	13.1	1.0a	27.0	1290	14.8	0 a	27.5
1241	13.1	1.0a	26.5	1291	14.2	T	27.5
1242	13.9	1.0a	27.0	1292	14.6	T	x
1243	14.2	1.0a	26.0	1293	13.8	T	x
1244	15.0	1.5a	25.0	1294	13.3	T	x
1245	15.5	2.0a	x	1295	14.4	T	26.5

Fix No.	Water depth	Postglacial sediment thickness	Till thickness	Fix No.	Water depth	Postglacial sediment thickness	Till thickness
1296 1297	14.7 14.8	T T	27.0 27.0	VERMILION			
1298	15.4	1.0a	25.0				
1299	15.5	1.5a	24.0	1	9.3	0	0
1300	15.7	1.5a	25.0	2	8.4	0	0
1301	15.4	1.0a	26.0	3	8.4	0	0
				4	8.3	0	0
1302			b	5	11.4	0	0
1303	14.9	T	x	7	10.6 11.6	0 0	0
1304	14.7	T	x	8	12.7	3.0	0 0
1305	13.4	1.0	x	9	12.7	5.5	0
1306	14.3	0.5	25.5	10	8.9	0	0
1307	14.5	1.0	24.0	11	9.5	5.5	0
1308	14.1	1.0	25.5	**	9.3	J.J	U
1309	13.4	1.0	25.5				
1310	14.0	0.5	25.5				
1311	15.1	1.5a	25.0				
1312	15.2	2.0a	23.5				
1313	15.0	3.0a	22.5				
1314	15.8	3.5a	22.5	1			
1315	15.4	4.0a	22.5				
1316	15.2	5.8	x				
1317	15.2	6.0	x				
1318	15.2	7.0	x				
1319	15.0	6.5	x				
1320	15.0	6.0	x				
1321	15.2	6.0	x				
1322	15.4	5.0	x				
1323	15.4	5.0	x				
1324	13.0	6.0	x				
1325	15.0	4.0	x				
1326	15.0	3.0	x				
1327	15.2	3.5	х	İ			
1328	15.0	7.5	x	}			
1329	15.0	8.0	х				
1330	15.0	6.8	x				
1331	15.0	6.0	x	•			
1332	14.5	5.0	X O E				
1333	14.0	4.5	8.5	}			
1334	13.2	2.5	8.5				
1335	13.0	1.0	8.0	1			
1336	12.2	1.0	6.0	1			
1337	11.8	1.0	3.0	ĺ			
1338	11.2	1.0	0	1			

APPENDIX G

CALCULATIONS OF SEISMIC VELOCITY IN WATER

Core No.	Measured depth	Seismic Time	Speed	Core No.	Measured depth	Seismic Time	Speed
	(m)	(ms)	(m/ms)		(m)	(ms)	(m/ms)
50	16.0	21.177	1.511	71	16.0	20.962	1.526
51	14.0	18.676	1.499	72	13.6	17.386	1.564
53	13.4	19.118	1.401	75	15.8	20.909	1.511
54	15.2	19.706	1.542	76	10.7	13.750	1.556
55	18.0	22.647	1.589	77	15.6	20.682	1.508
56	15.9	20.295	1.566	78	15.4	20.228	1.522
58	10.4	13.382	1.554	80	10.5	13.296	1.579
59	12.0	14.707	1.631	82	9.1	12.159	1.496
60	13.0	16.764	1.550	84	15.2	19.772	1.537
61	16.0	21.029	1.521	86	15.8	20.439	1.546
62	17.6	22.500	1.564	87	15.3	19.863	1.540
64	15.1	19.176	1.574	89	13.4	17.472	1.533
67	20.7	26.591	1.556	92	13.9	17.955	1.548
69	18.3	23.294	1.571	97	13.3	17.143	1.551
70	11.3	13.863	1.630	100	8.9	12.273	1.450

APPENDIX H

CALCULATIONS OF SEISMIC VELOCITY IN POSTGLACIAL SEDIMENTS

Core No.	Postglacial sediment thickness (m)	Time to Pleistocene reflector (ms)	Seismic speed (m/ms)
54	1.41	3.00	0.94
55	3.90	5.50	1.42
60	5.57	6.24	1.78
61	3.05	6.00	1.02
77	1.54	2.50	1.23
78	1.25	2.00	1.25
79	0.80	1.50	1.07
84	2.98	4.00	1.49
85	3.90	5.00	1.56
97	1.02	1.50	1.60
98	3.03	2.00	1.21
99	2.18	3.50	1.25

APPENDIX I

COMPARISON OF POSTGLACIAL SEDIMENT THICKNESSES
AS MEASURED FROM THE CORES AND CALCULATED FROM THE SEISMIC RECORDS

Core		Thickness	
No	Seismic (m)	Measured (m)	Difference (m)
51	2.5	3.7	1.2
54	2.0	1.4	-0.6
55	3.5	3.9	0.4
60	4.0	5.6	1.6
61	4.0	3.0	-1.0
73	0	0.3	0.3
75	0	0.2	0.2
76	1.0	0.6	-0.4
77	1.5	1.5	0.0
78	1.5	1.3	-0.2
79	1.0	0.8	-0.2
81	3.5	3.0	-0.5
84	2.5	3.0	0.5
85	3.5	3.8	0.3
86	2.0	2.1	0.1
92	1.0	1.1	0.1
95	4.0	4.0	0.0
96	3.0	2.5	-0.5
97	1.0	1.0	0.0
99	2.5	2.2	-0.3

Corps of Engineers, Coastal Engineering Research Center, Springfield, Lake Erie. The primary objective is to locate and delineate sand and Corps of Engineers, Coastal Engineering Research Center, Springfield, [109] p.: 111.; 28 cm.--(Miscellaneous report / Coastal Engineer-ing Research Center; no. 82-15). Cover title.
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